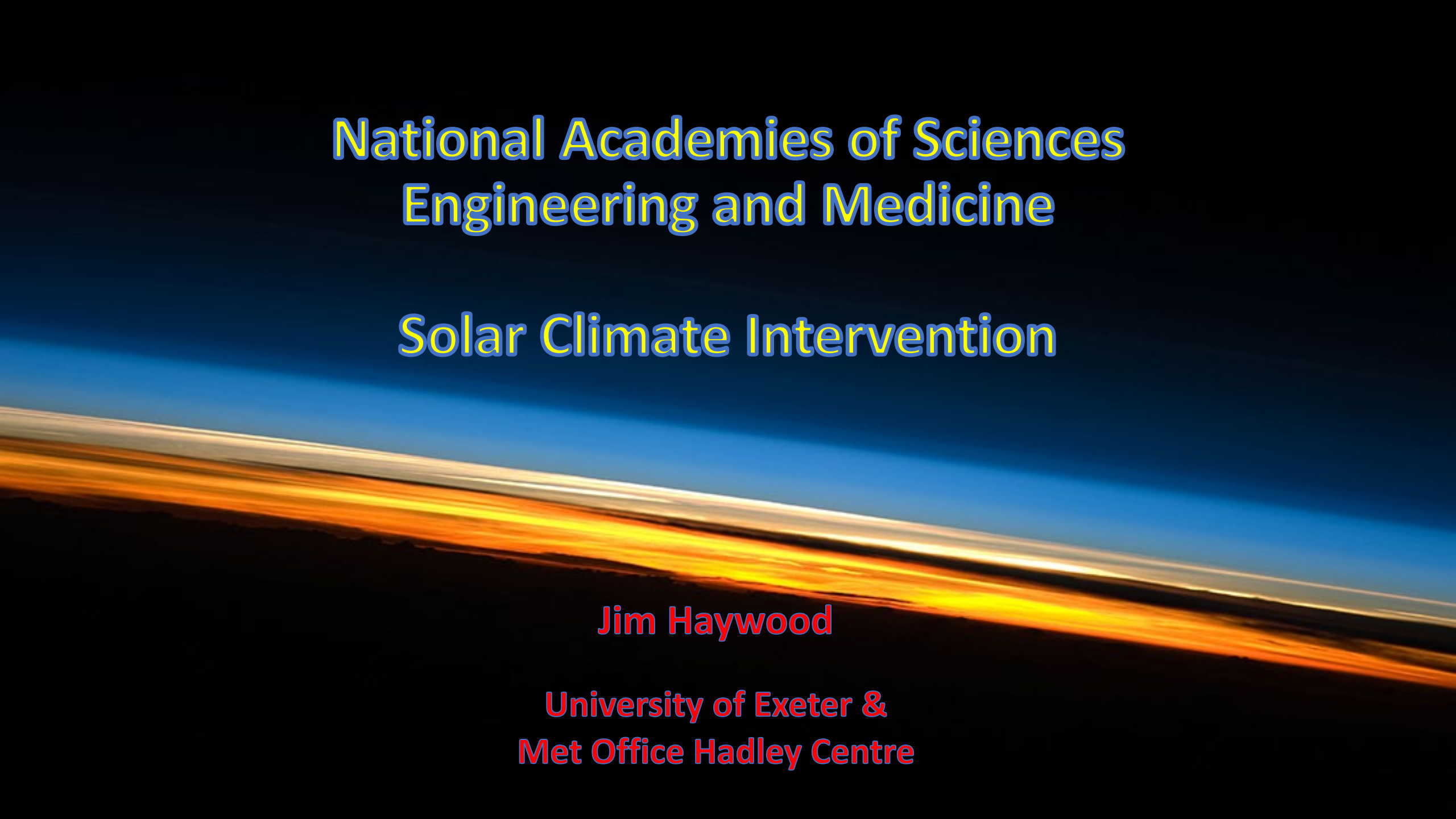




National Academies of Sciences Engineering and Medicine

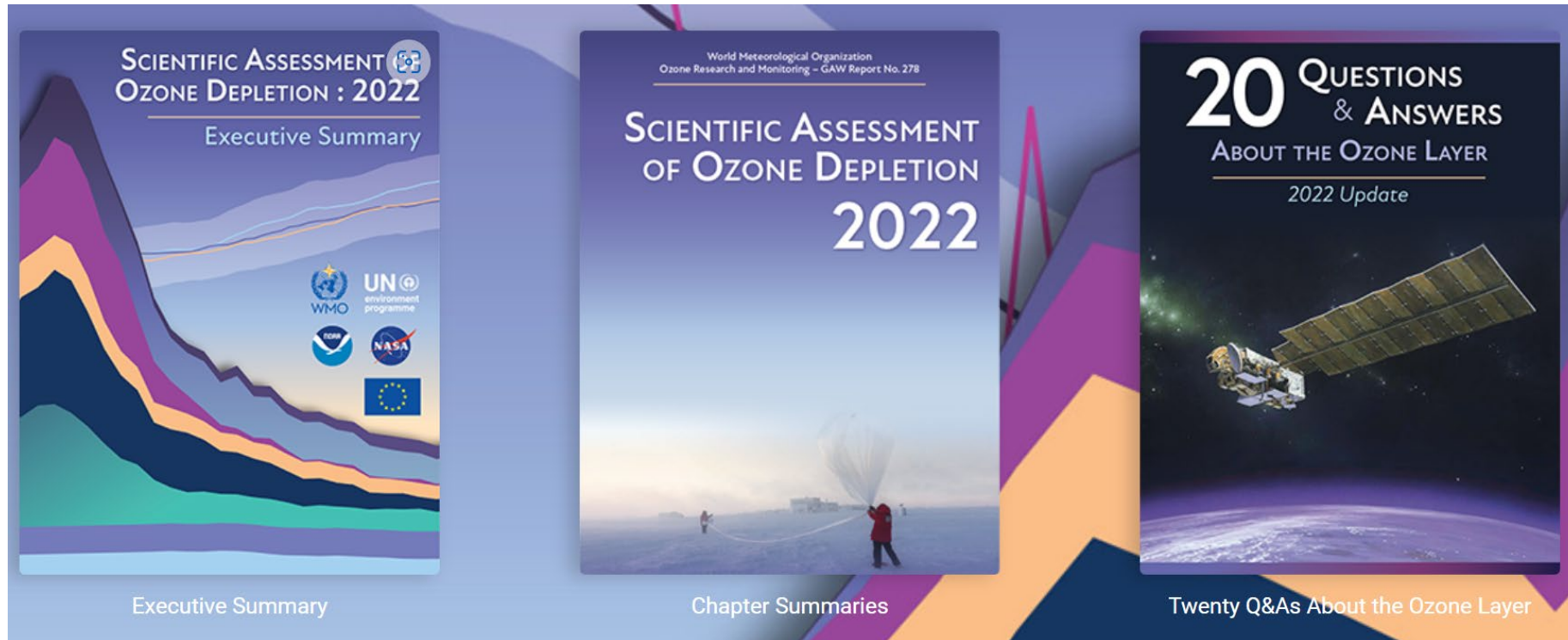
Solar Climate Intervention

Jim Haywood

University of Exeter &
Met Office Hadley Centre

Solar Climate Intervention is being taken seriously by Policy Makers.

The WMO/United Nations have included deliberate stratospheric aerosol injection in their ozone depletion assessment for the first time.

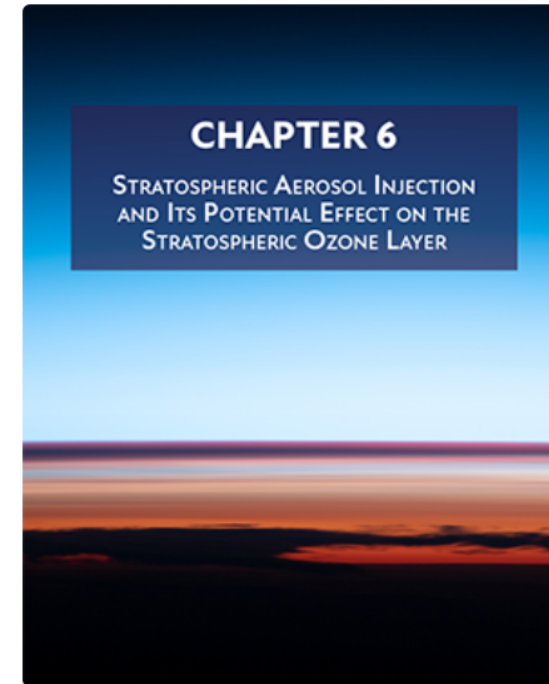


<https://csl.noaa.gov/assessments/ozone/2022/>

The rationale for this technology is clear

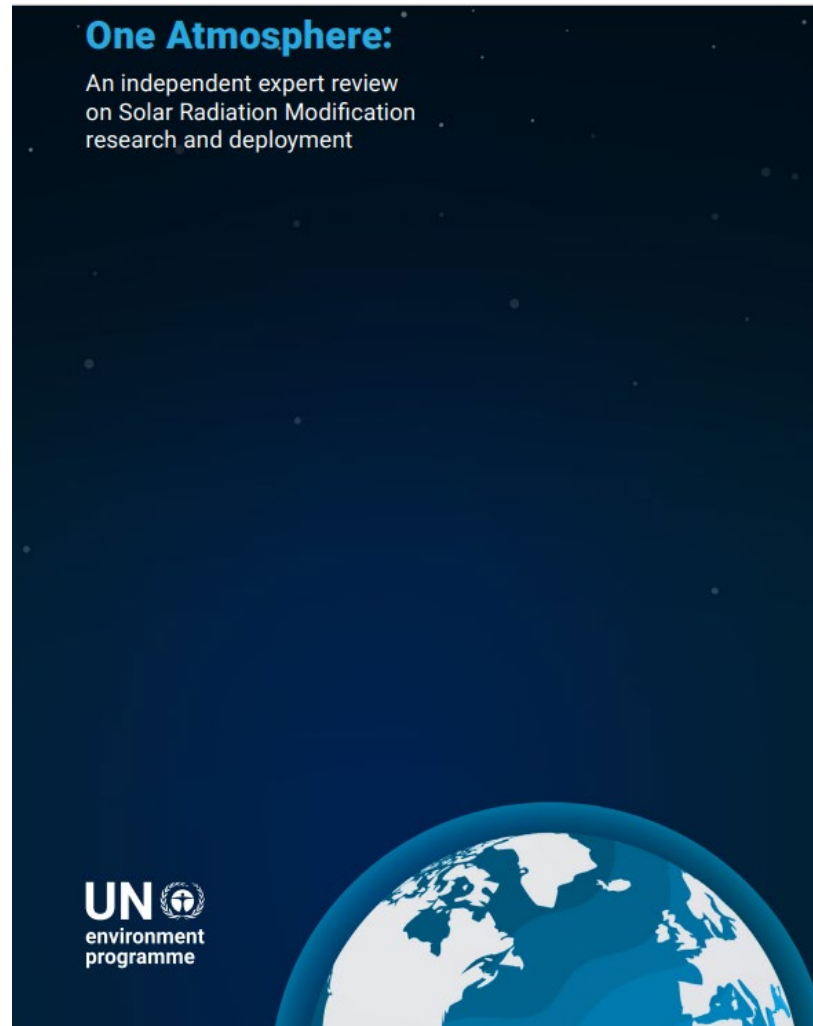
Chapter 6: Stratospheric Aerosol Injection and Its Potential Effect on the Stratospheric Ozone Layer

Since the 2018 Ozone Assessment global warming has continued, having now reached approximately 1.2°C above preindustrial levels. All climate model scenarios considered by IPCC (2021) indicate continued future warming beyond 1.5°C above the preindustrial level, a limit that has been proposed to prevent further detrimental impacts. Ambitious mitigation and decarbonization efforts are required to minimize the likely overshoot of temperatures above this limit and to stabilize global surface temperatures in the future. However, with a temperature overshoot, irreversible impacts on the climate system may still occur. Stratospheric aerosol injection (SAI) has been suggested as a potential mechanism for reflecting sunlight back to space, thereby offsetting some of the surface warming. Evidence from explosive volcanic eruptions and various model simulations has shown that increasing stratospheric sulfate aerosols can substantially cool the planet. SAI and other solar radiation modification (SRM) approaches may therefore be the only option to keep the global surface temperature below the limit of 1.5°C. The amount and duration of SAI required would depend on how fast atmospheric greenhouse gas (GHG) concentrations are lowered through mitigation and decarbonization efforts.



- 1) We're going to overshoot 1.5C
- 2) Our mitigation efforts are not enough to prevent this
- 3) Stratospheric Aerosol Injection may be the only option available to us to avoid hazardous impacts

The UN have commissioned and endorsed independent expert review



Foreword



Make no mistake: there are no quick fixes to the climate crisis. Increased and urgent action to slash greenhouse gas emissions and invest in adapting to the impacts of climate change is immutable. Yet current efforts remain insufficient. As a result, increasing voices are calling for and preparing alternative "emergency" options to keep global temperature rise in check.

Among actions under examination is Solar Radiation Modification (SRM), and in particular Stratospheric Aerosol Injection (SAI) – which aims to cool the planet by reflecting sunlight back into space. SRM is a complex, controversial and under-studied group of technologies. Yet some scientists and companies are accelerating towards deployment: empirical research and experimentation are being pursued, and technologies and schemes are being discussed at the highest levels, without a full understanding of the implications. This is contrary to the precautionary principle, which must be applied in the case of a technology that would modify the atmosphere.

<https://www.unep.org/resources/report/Solar-Radiation-Modification-research-deployment>



SPACE MIRRORS

Orbiting mirrors deflect sun's rays

READINESS: ●●●●●

COST: \$\$\$

FLAW: unknown weather effects; fails to prevent acidic oceans



ARTIFICIAL TREES

CO₂ sucked from air and stored underground

READINESS: ●●●●●

COST: \$\$\$

FLAW: large geological cache needed



REFLECTIVE CROPS

Planting crops that reflect more sunlight

READINESS: ●●●●●

COST: \$

FLAW: large land area needed; fails to prevent acidic oceans



FORESTING

Trees absorb CO₂

READINESS: ●●●●●

COST: \$

FLAW: large land area needed



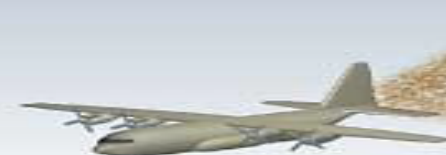
AEROSOLS

Particles in the stratosphere reflect sun's rays

READINESS: ●●●●●

COST: \$

FLAW: risk of ozone depletion; unknown weather effects; fails to prevent acidic oceans



CLOUD SEEDING

Atomising seawater creates clouds to reflect sun's rays

READINESS: ●●●●●

COST: \$\$

FLAW: unknown weather effects; patchy success; fails to prevent acidic oceans



BIOCHAR

Agricultural carbon waste is burned and buried

READINESS: ●●●●●

COST: \$\$

FLAW: large land area needed



OCEAN FERTILISATION

Iron filings stimulate CO₂-eating plankton

READINESS: ●●●●●

COST: \$\$

FLAW: unknown effects on ecosystems



CARBONATE ADDITION

Ground limestone helps oceans absorb CO₂

READINESS: ●●●●●

COST: \$\$

FLAW: unknown effects on ecosystems

● Cooling factor: potential to change Earth's energy budget

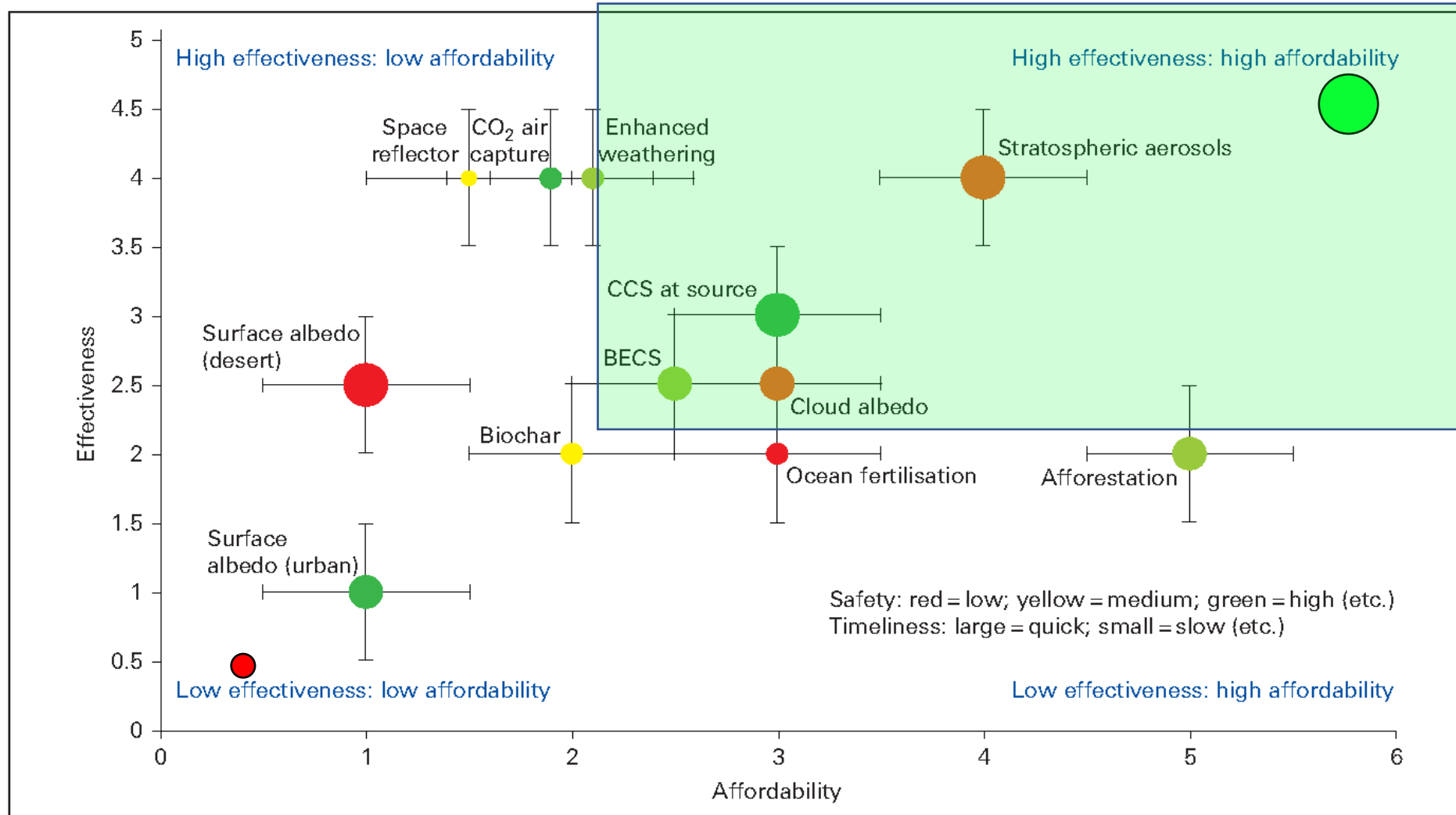
Readiness:

● - Within years
●● - Within decades
●●● - Within centuries

Cost:

\$ - Cheap relative to cutting emissions
\$\$ - Significant compared to cost of cutting emissions
\$\$\$ - Cutting emissions might be cheaper

New Scientist, 2009





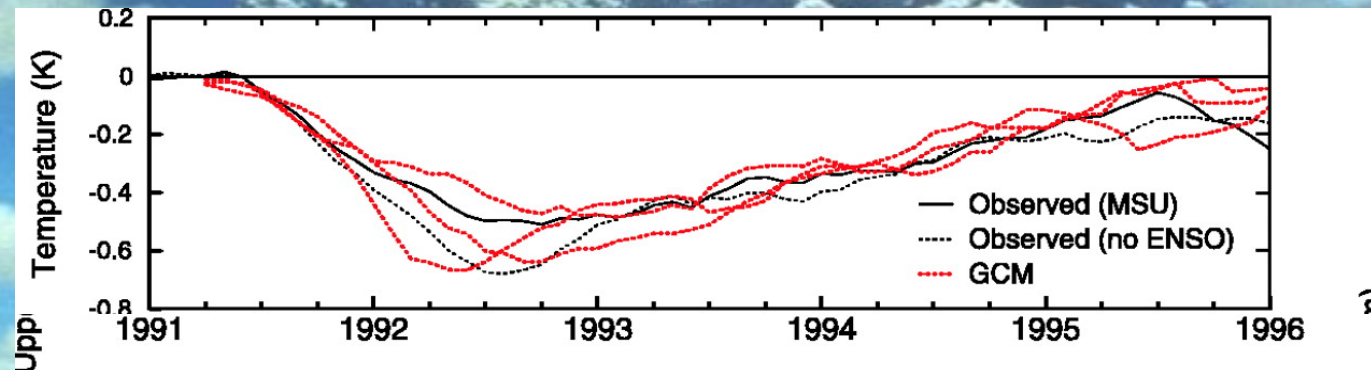
Stratospheric Aerosol Injection (SAI)



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of Exeter

- The most discussed and most widely modelled SRM option.
- Mimics the cooling impact from volcanoes
- Majority of studies suggest injection of SO₂
- Potential delivery mechanisms include high altitude aircraft, tethered balloons, rockets, artillery.

Mt. Pinatubo, 1991



Volcanoes cause global cooling by putting aerosols in the stratosphere: aerosols last for 1-2 years in the stratosphere (e.g. Soden et al., 2022).

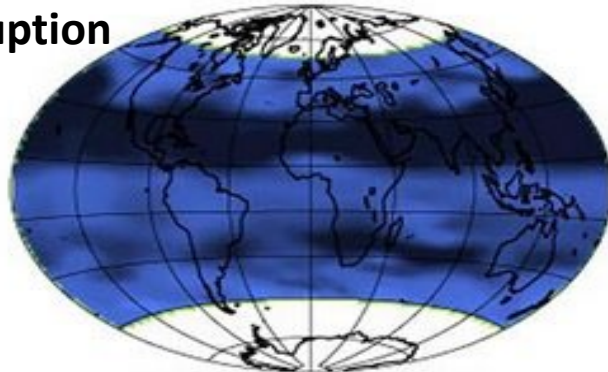
There is debate about the degree of cooling in the observational record owing to concurrent climate variability.



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SAGE II 1020 nm optical depth

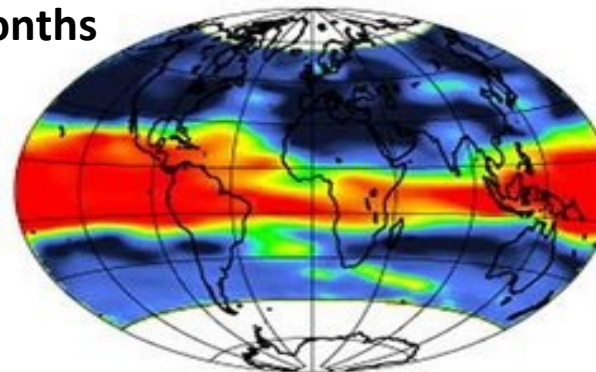
Pre-eruption



April 10, 1991 to May 13, 1991

(a)

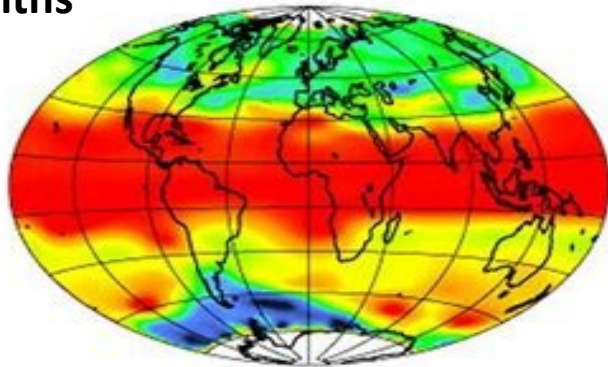
+3 months



June 15, 1991 to July 25, 1991

(b)

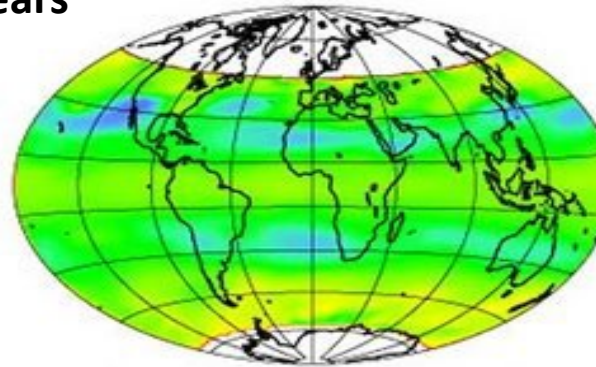
+5 months



August 23, 1991 to September 30, 1991

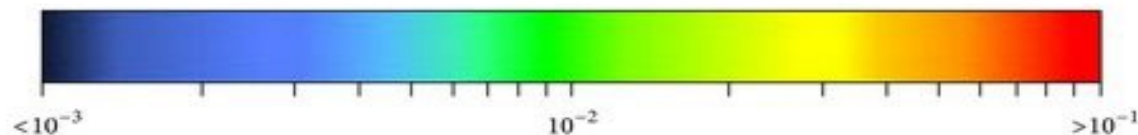
(c)

+2 years



December 5, 1993 to January 16, 1994

(d)

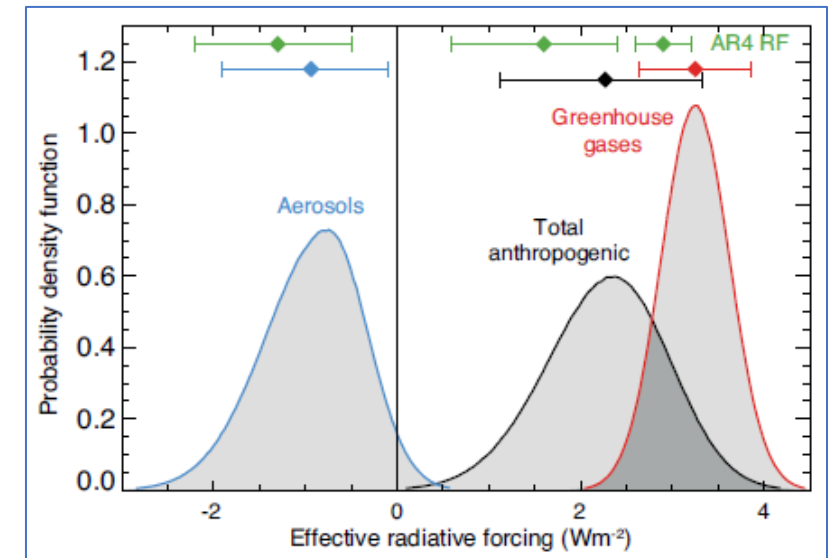


Aerosol Optical Depth
observations from the
Stratospheric Aerosol and Gas
Experiment (SAGE II)

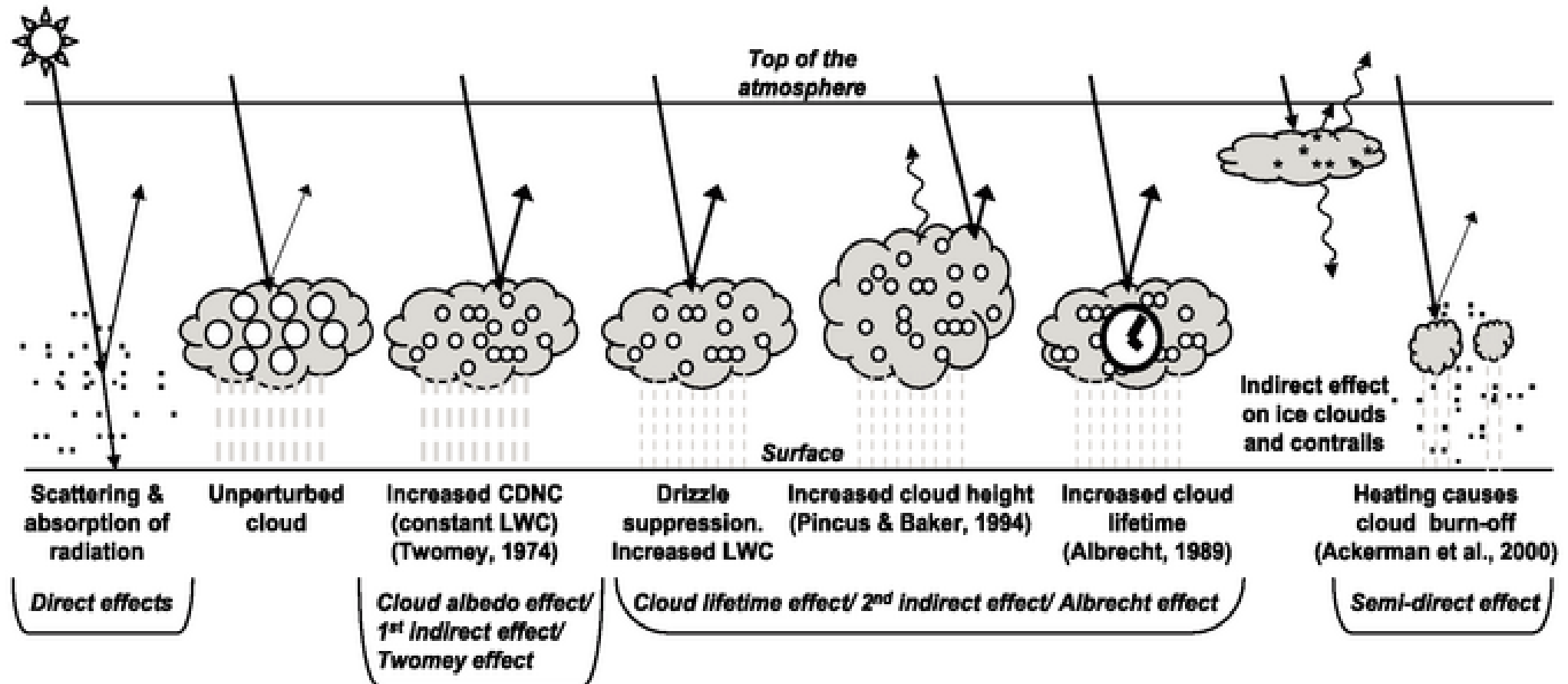
Aerosol is transported from the
equator to the poles in the
stratospheric Brewer-Dobson
circulation evolving into a more
uniform distribution

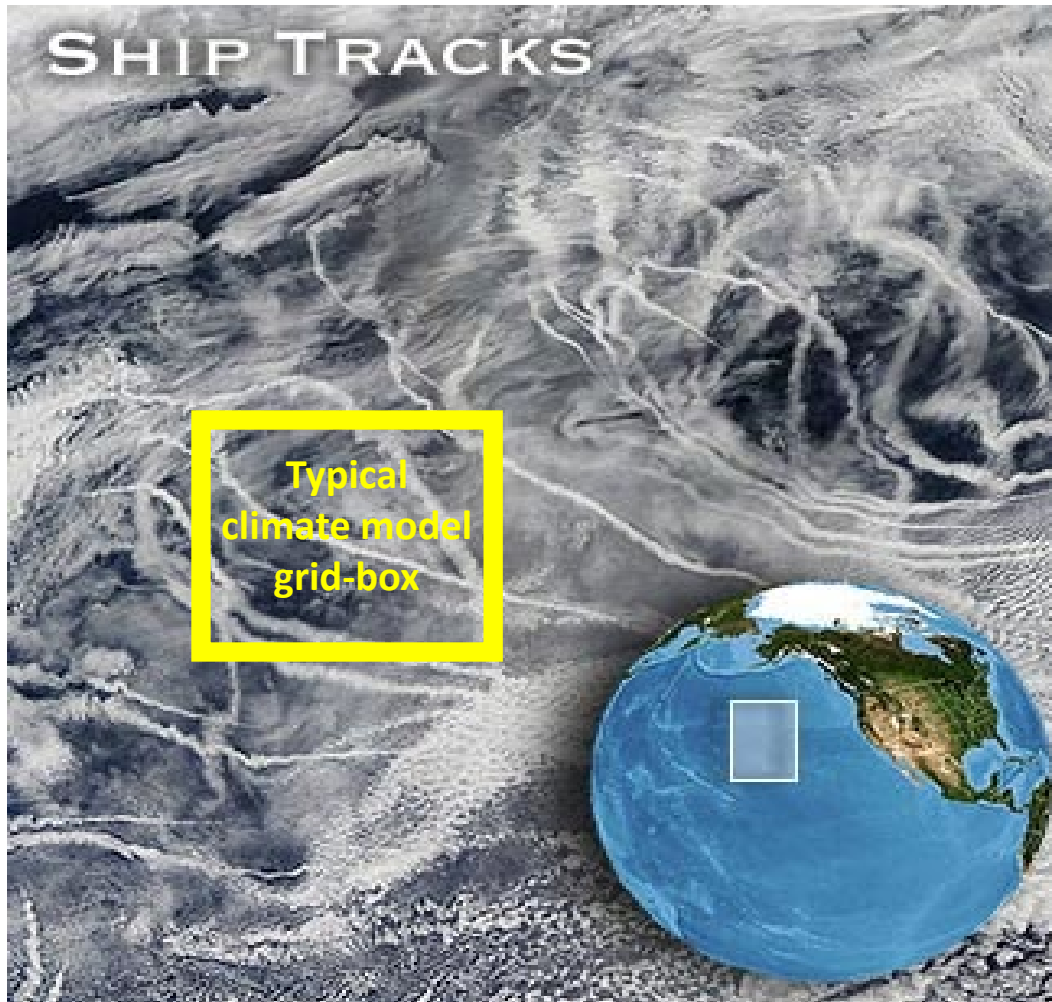
Marine Cloud Brightening (MCB)

- Evidence of MCB frequently from ship tracks
- More large scale evidence of MCB from massive effusive volcanic eruptions
- Majority of studies suggest injection of sea-salt
- Aerosol-cloud interactions are generally poorly constrained in global climate models



Marine Cloud Brightening





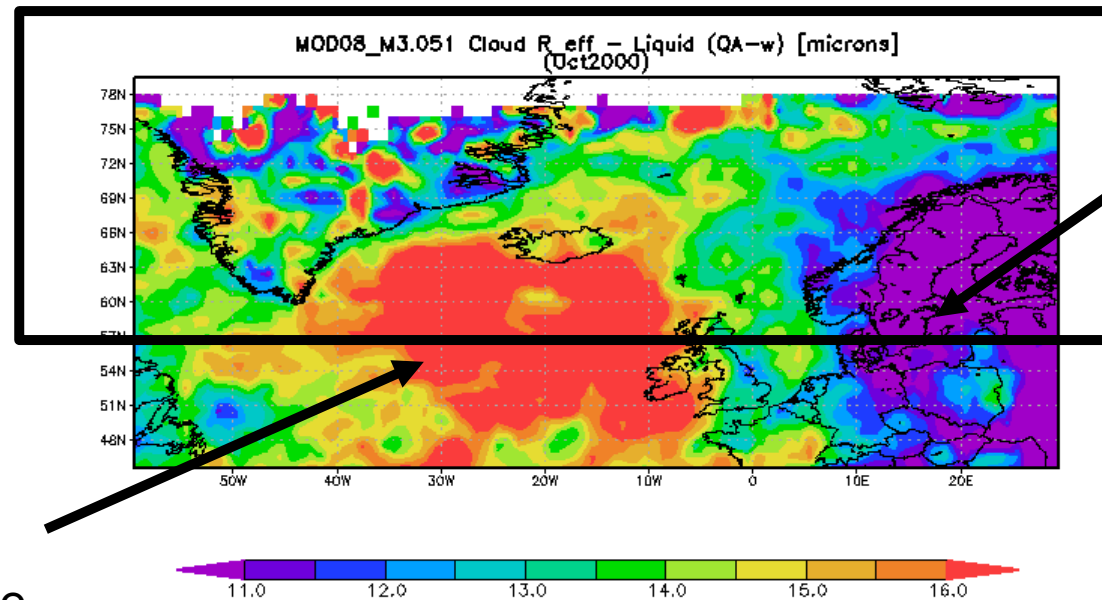
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To challenge global climate models we need perturbations that are:

- **large-scale**
- **long-lived**
- **in pristine environments**

There is clear observational evidence for aerosol-cloud interactions from ship-tracks, but these features are sub-gridscale when compared to the resolution of typical global climate models (GCMs).

2000

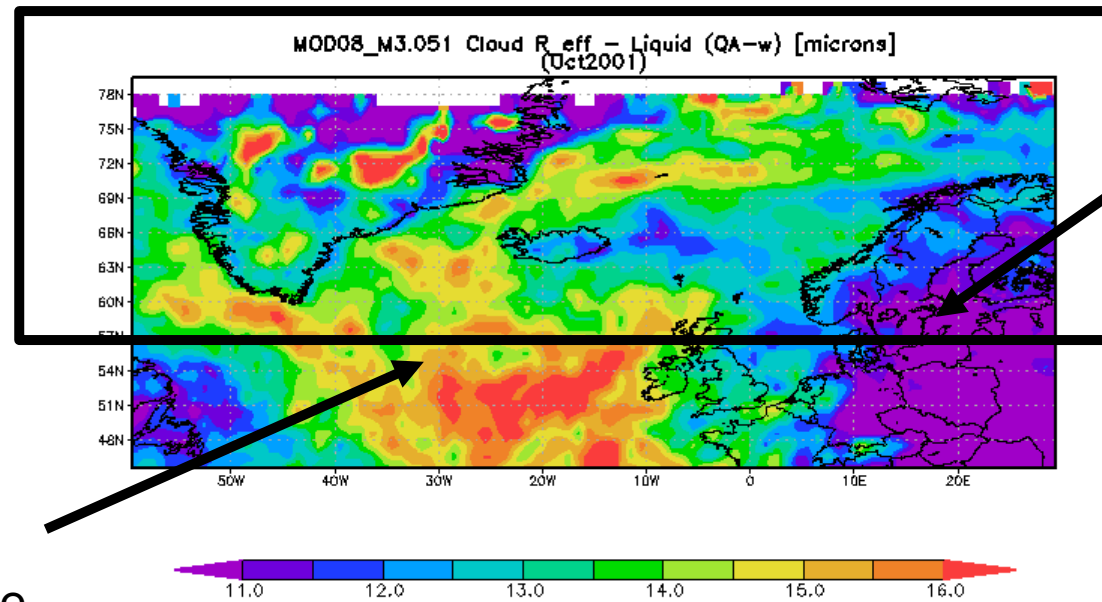


Marine clouds:
Cloud top effective
radius ~15microns

Continental clouds:
Cloud top effective
radius ~11microns

MODIS Terra and Aqua (22
year record) provide an
invaluable tool for studying
aerosol-cloud-interactions.

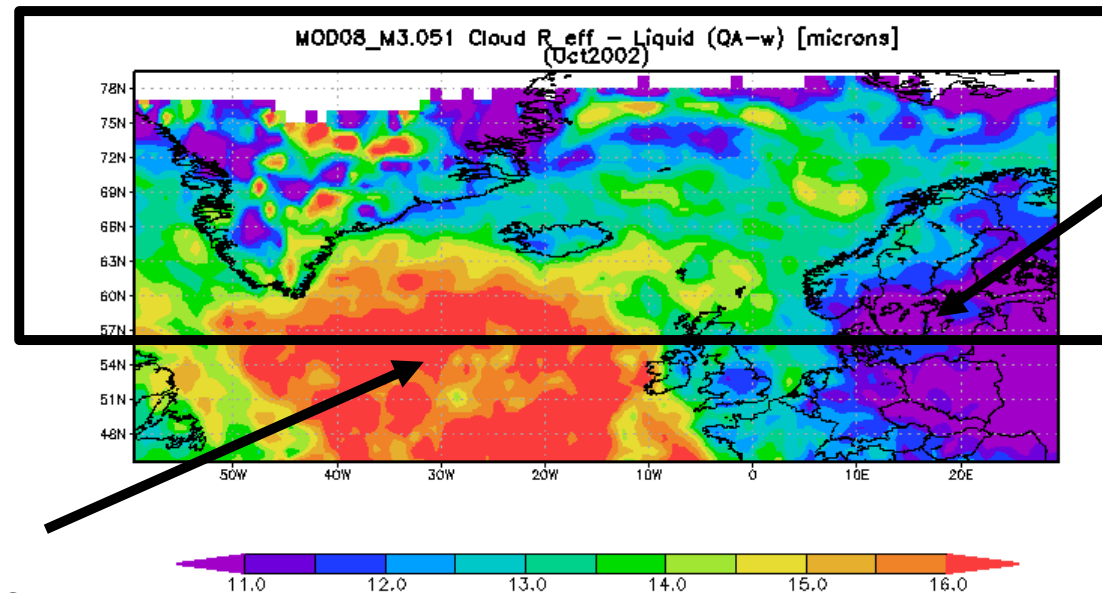
2001



Continental clouds:
Cloud top effective
radius ~11microns

Marine clouds:
Cloud top effective
radius ~15microns

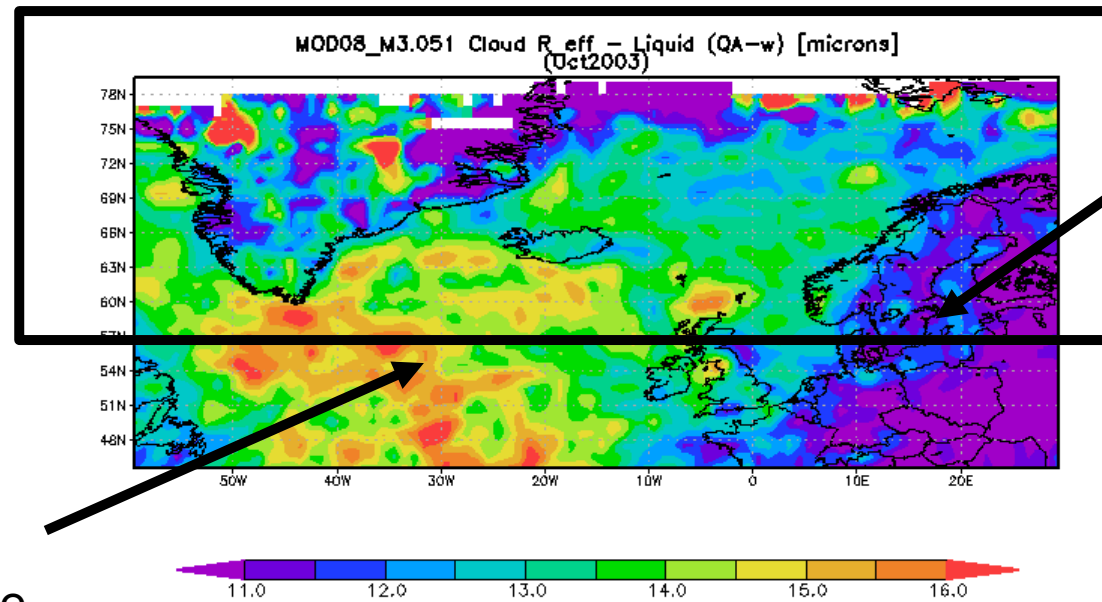
2002



Continental clouds:
Cloud top effective
radius ~11microns

Marine clouds:
Cloud top effective
radius ~15microns

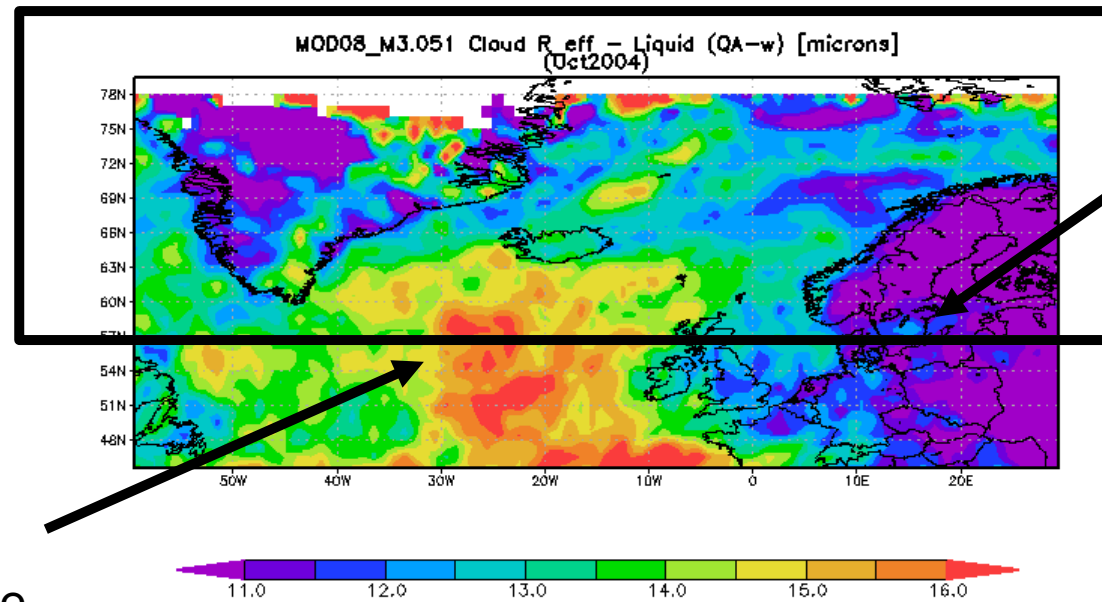
2003



Continental clouds:
Cloud top effective
radius ~11microns

Marine clouds:
Cloud top effective
radius ~15microns

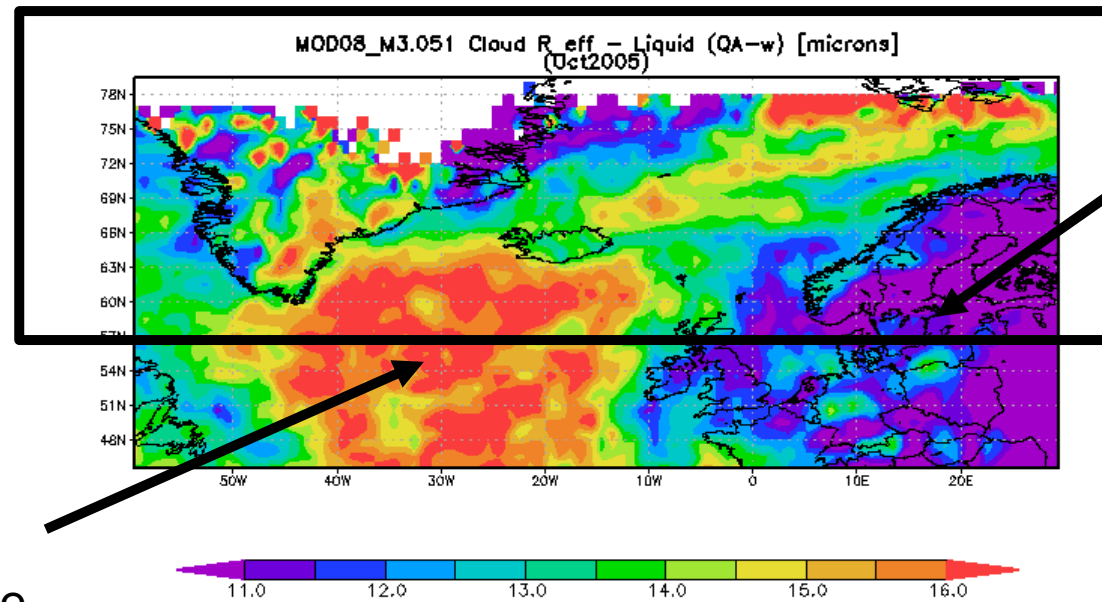
2004



Continental clouds:
Cloud top effective
radius ~11microns

Marine clouds:
Cloud top effective
radius ~15microns

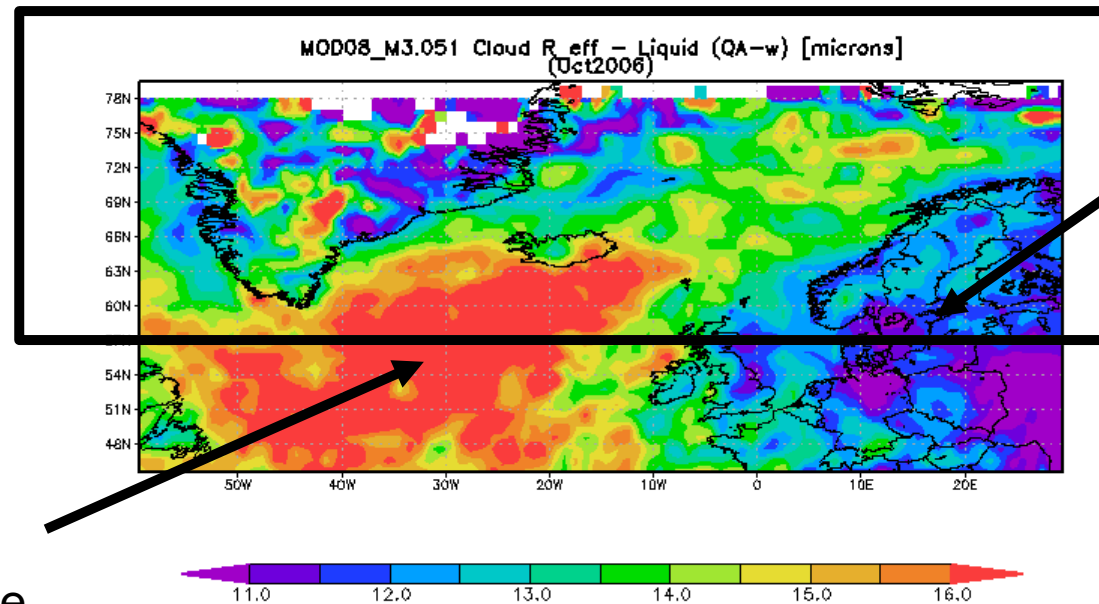
2005



Continental clouds:
Cloud top effective
radius ~11microns

Marine clouds:
Cloud top effective
radius ~15microns

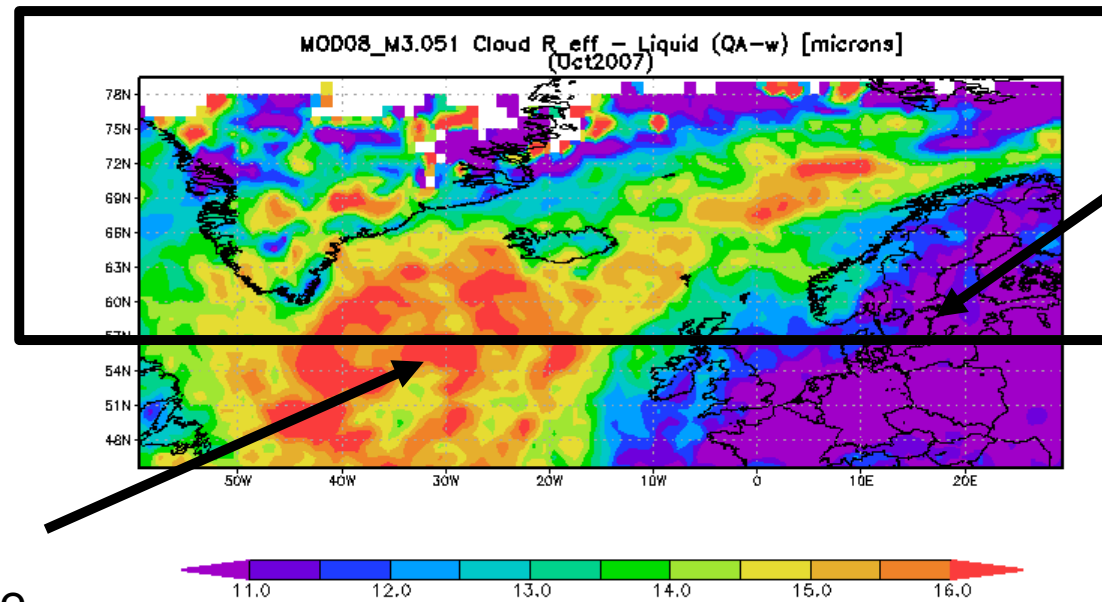
2006



Continental clouds:
Cloud top effective
radius ~11microns

Marine clouds:
Cloud top effective
radius ~15microns

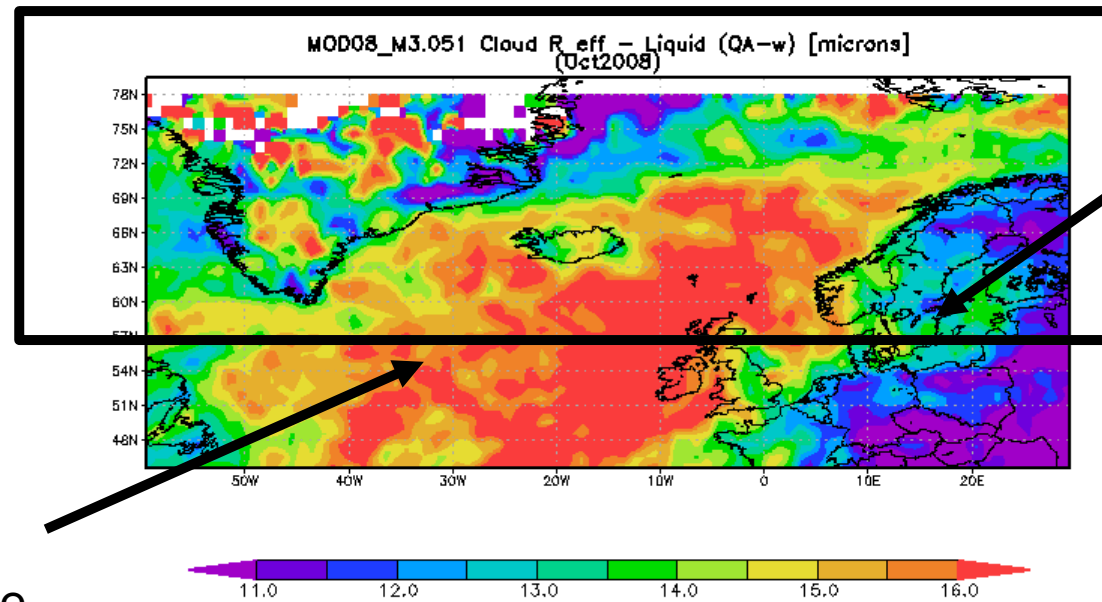
2007



Continental clouds:
Cloud top effective
radius ~11microns

Marine clouds:
Cloud top effective
radius ~15microns

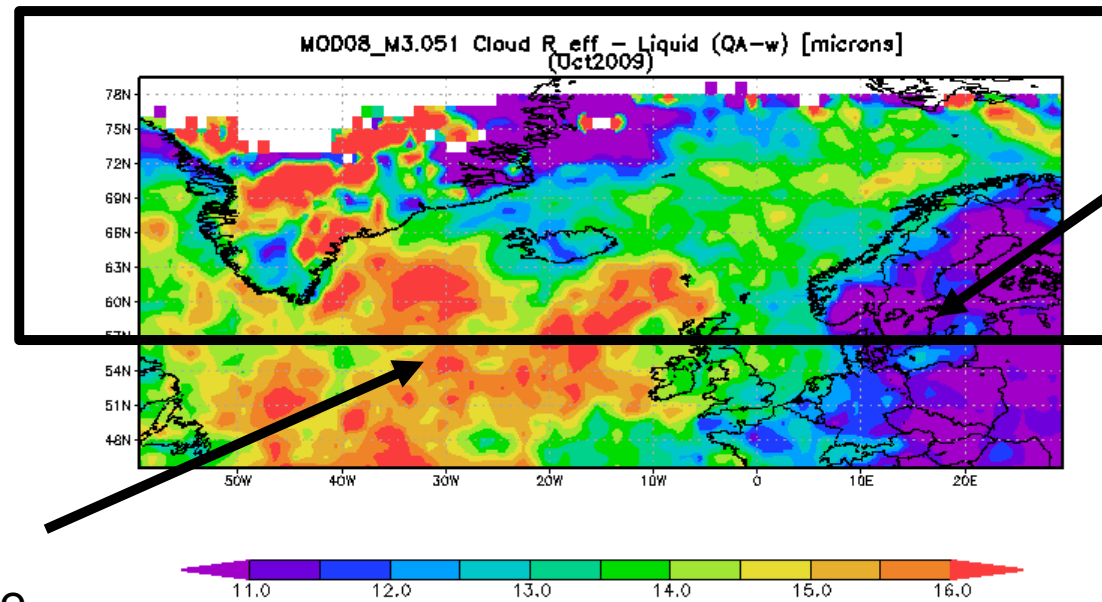
2008



Continental clouds:
Cloud top effective
radius ~11microns

Marine clouds:
Cloud top effective
radius ~15microns

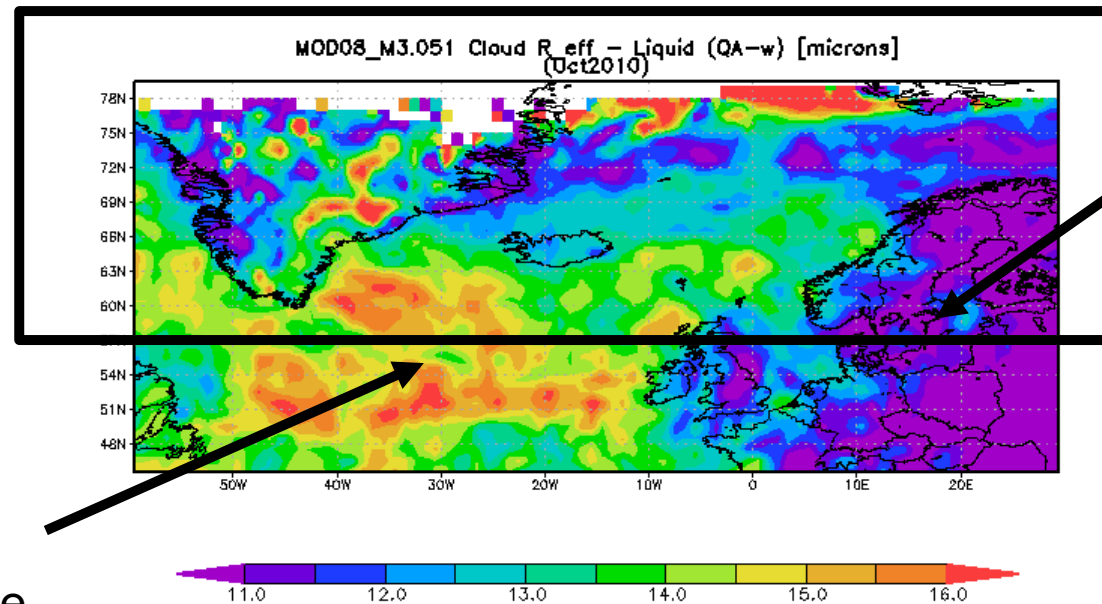
2009



Continental clouds:
Cloud top effective
radius ~11microns

Marine clouds:
Cloud top effective
radius ~15microns

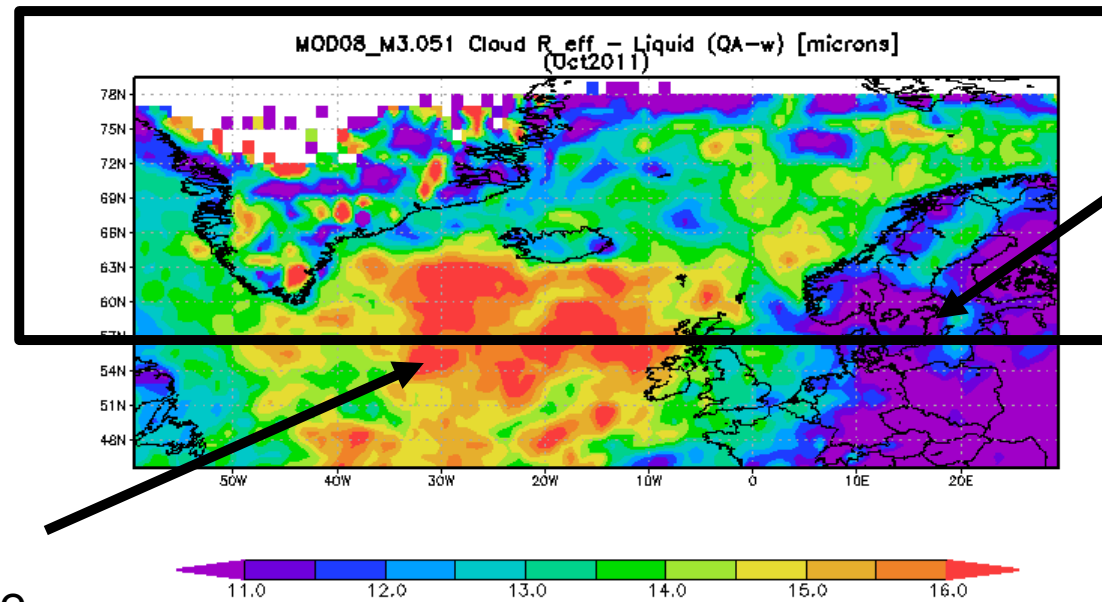
2010



Continental clouds:
Cloud top effective
radius ~11microns

Marine clouds:
Cloud top effective
radius ~15microns

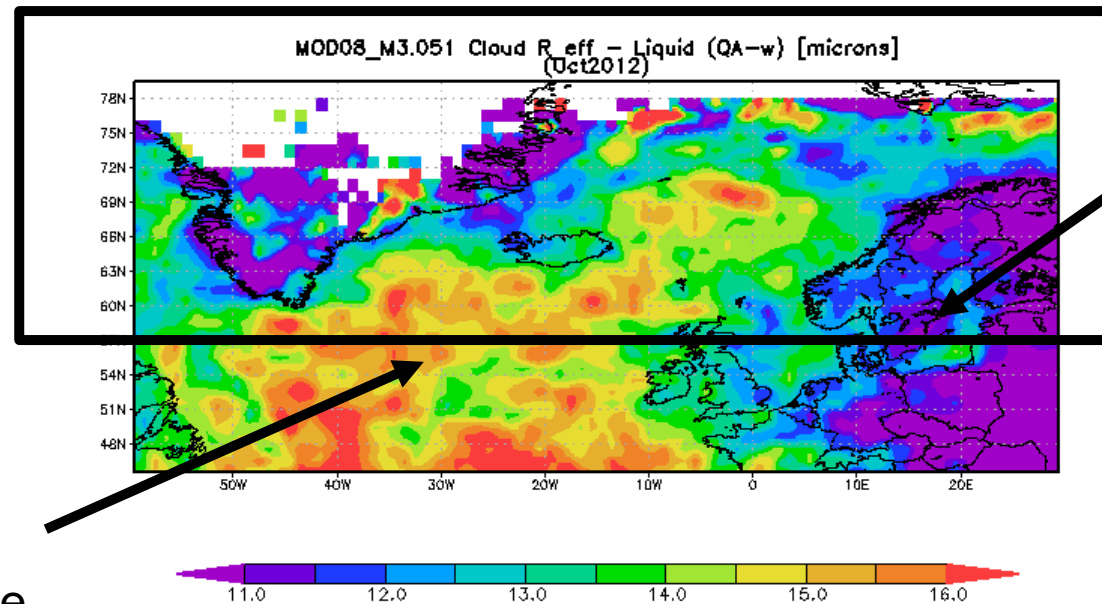
2011



Continental clouds:
Cloud top effective
radius ~11microns

Marine clouds:
Cloud top effective
radius ~15microns

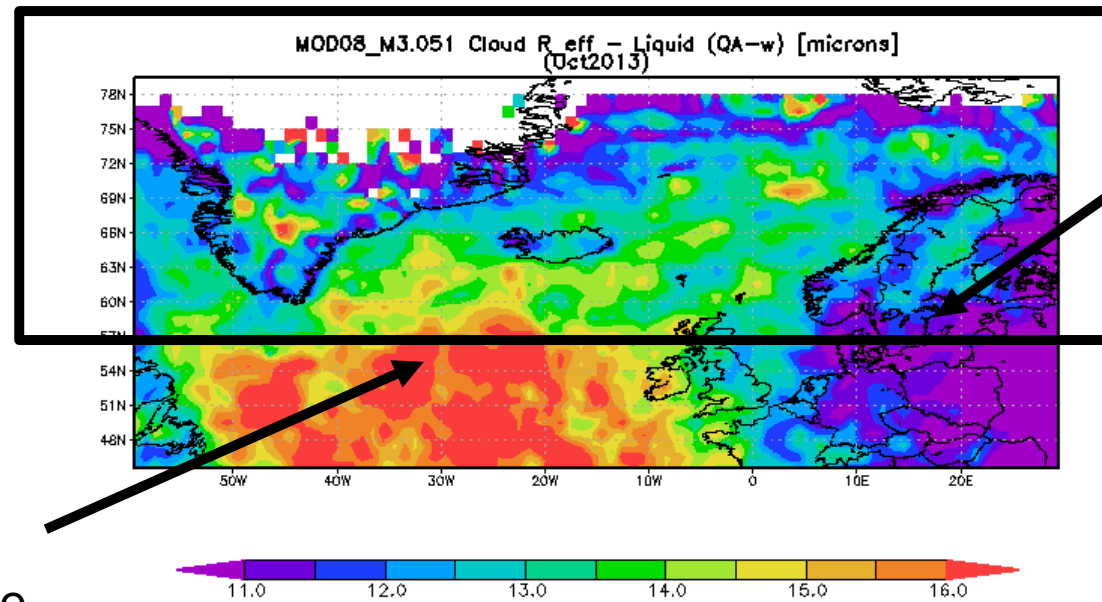
2012



Continental clouds:
Cloud top effective
radius ~11microns

Marine clouds:
Cloud top effective
radius ~15microns

2013

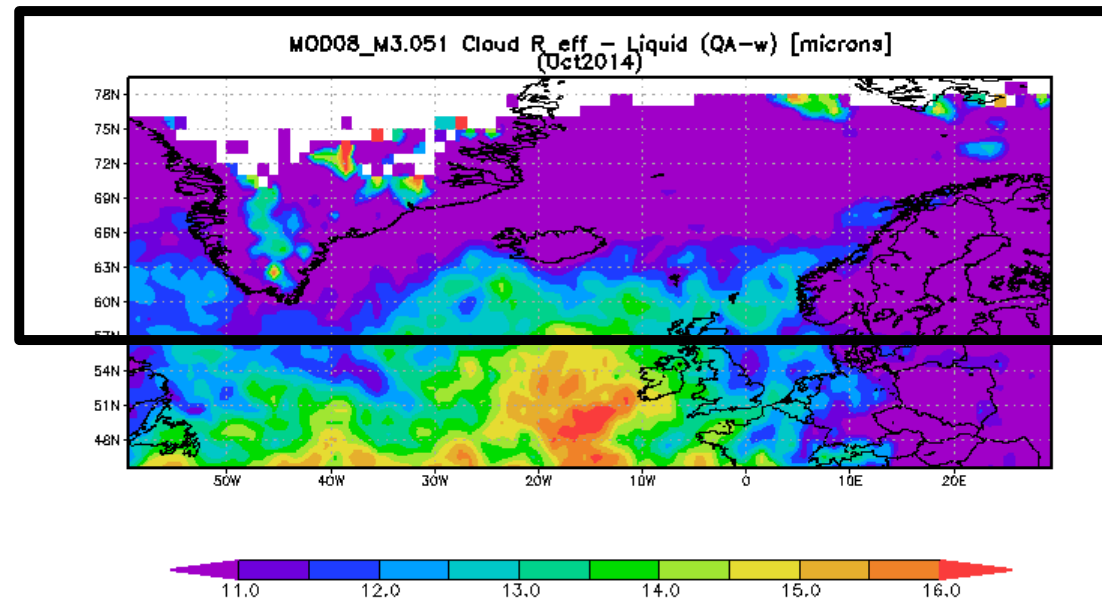


Continental clouds:
Cloud top effective
radius ~11microns

Marine clouds:
Cloud top effective
radius ~15microns

2014

Something
has
happened!



Over vast swathes of the N Atlantic:

“Dark Marine Clouds -> Bright
Continental clouds”

Holuhraun fissure eruption:

85km² lava field

Up to 100ktSO₂/day

Over x10 emission rate from all of 28 European countries put together (1/3 of global emissions of SO₂!)

Sustained for ~ 6 months

Image courtesy of Anja Schmidt





Cirrus Cloud Thinning



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- The least well studied
- The least well modelled
- Cirrus clouds impact the terrestrial radiation more than solar radiation
- How thinning could be achieved is an open question

‘Cirrus Cloud Thinning’

Thin cirrus cloud warm
the climate.

It crosses over a little
with contrail avoidance.

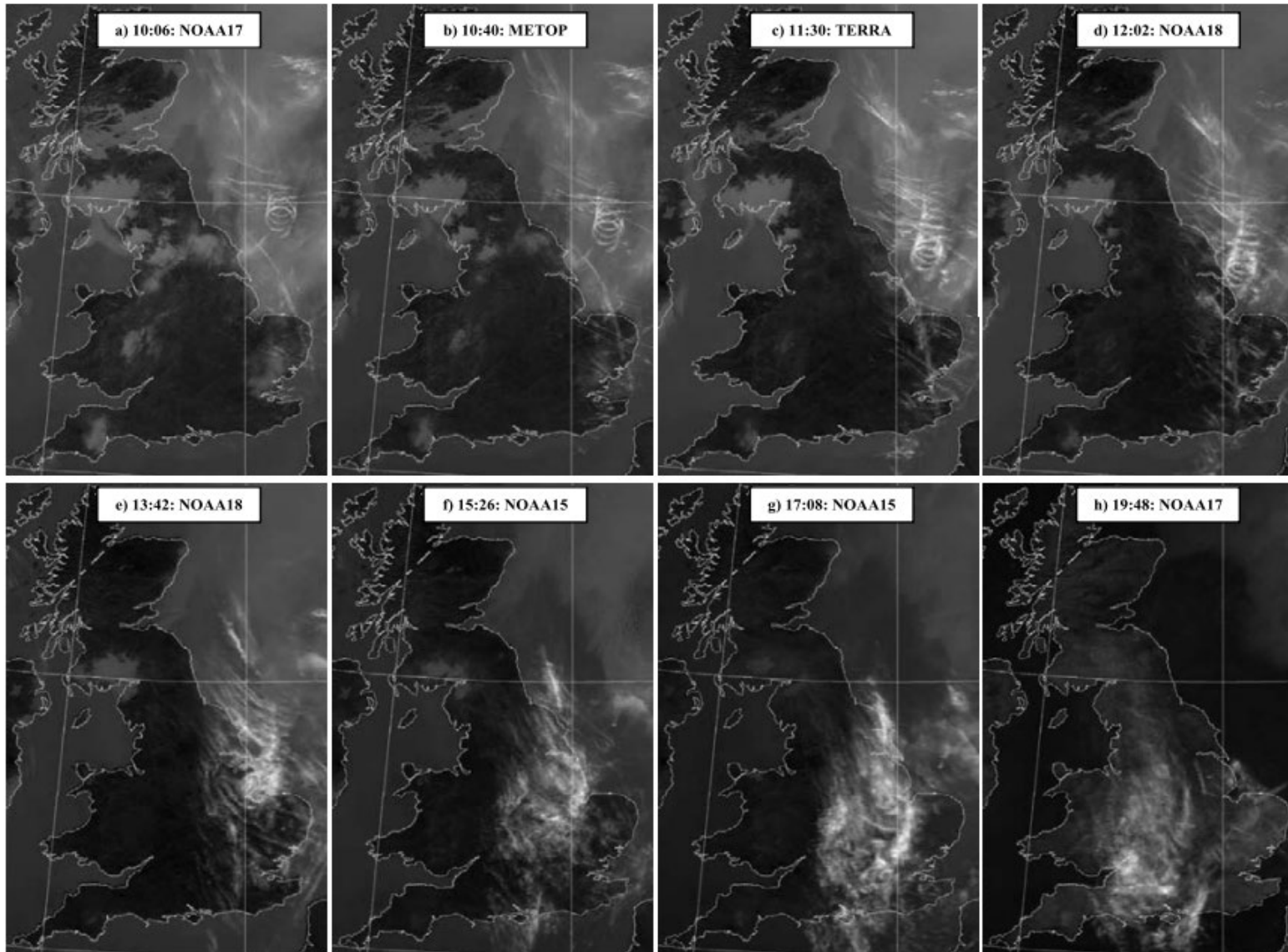


Figure 5. IR ($10.8\ \mu\text{m}$) images of the formation of contrail induced cirrus (bright white). Areas of stratocumulus are shown as medium gray. The time and satellite is shown in the inset in each of the frames.



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SAI: The state of the science global modelling

Most coordinated modelling is through GeoMIP



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- <http://climate.envsci.rutgers.edu/GeoMIP/>



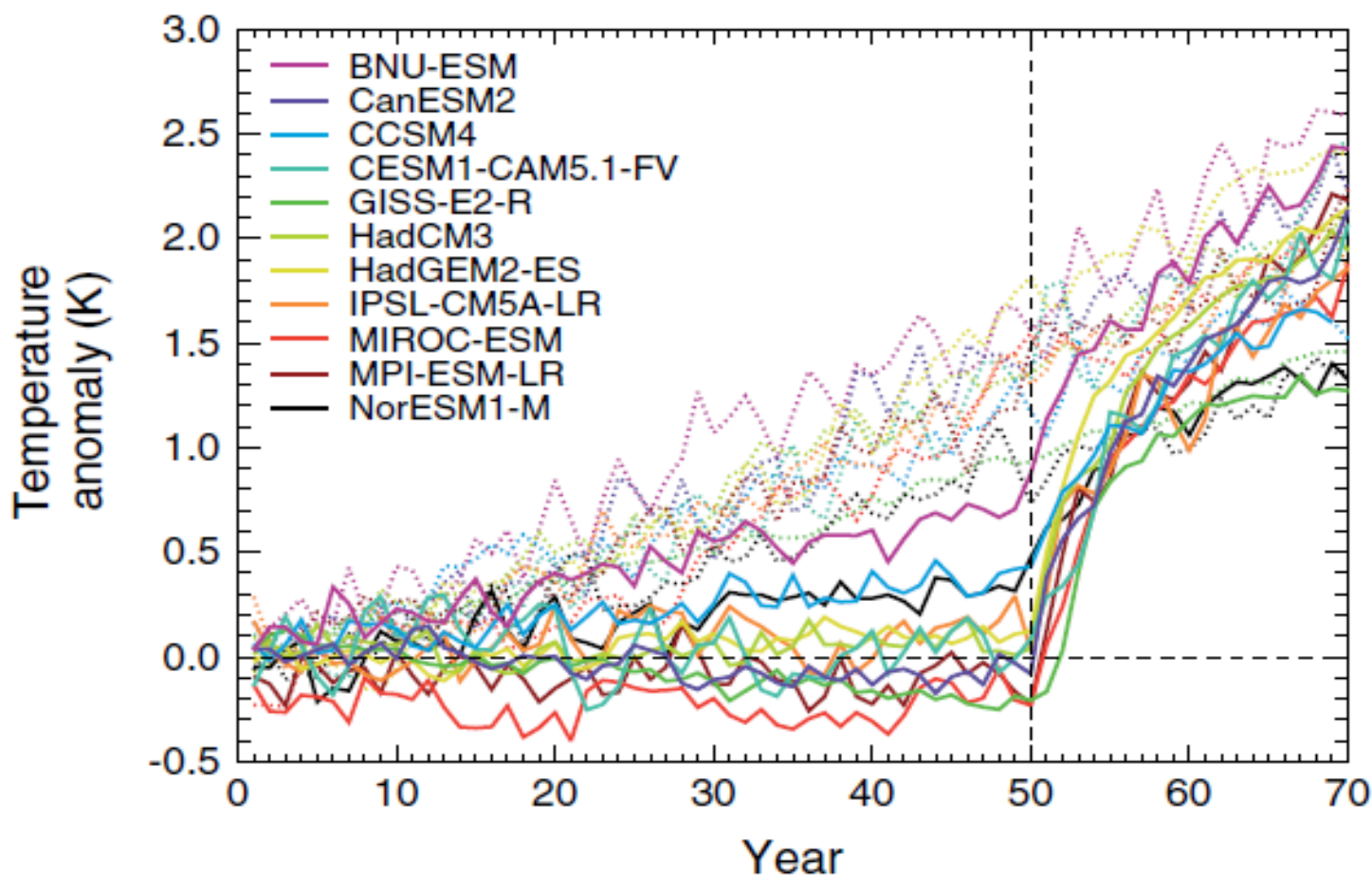
GeoMIP is self-funded rather than centrally funded

2023 meeting will have ~100 participants.

Early GeoMIP studies simply turned the sun down to balance a step GHG forcing so that all models could participate.

More recent studies have incorporated detailed sulphur aerosol microphysics and arguably more “realistic” scenarios

There is a balance between number of models and complexity



11 models participated in 'G2' – turning the sun down.

All models show 'termination effect'; all the avoided global warming is realised in a few years (Jones et al., 2013).

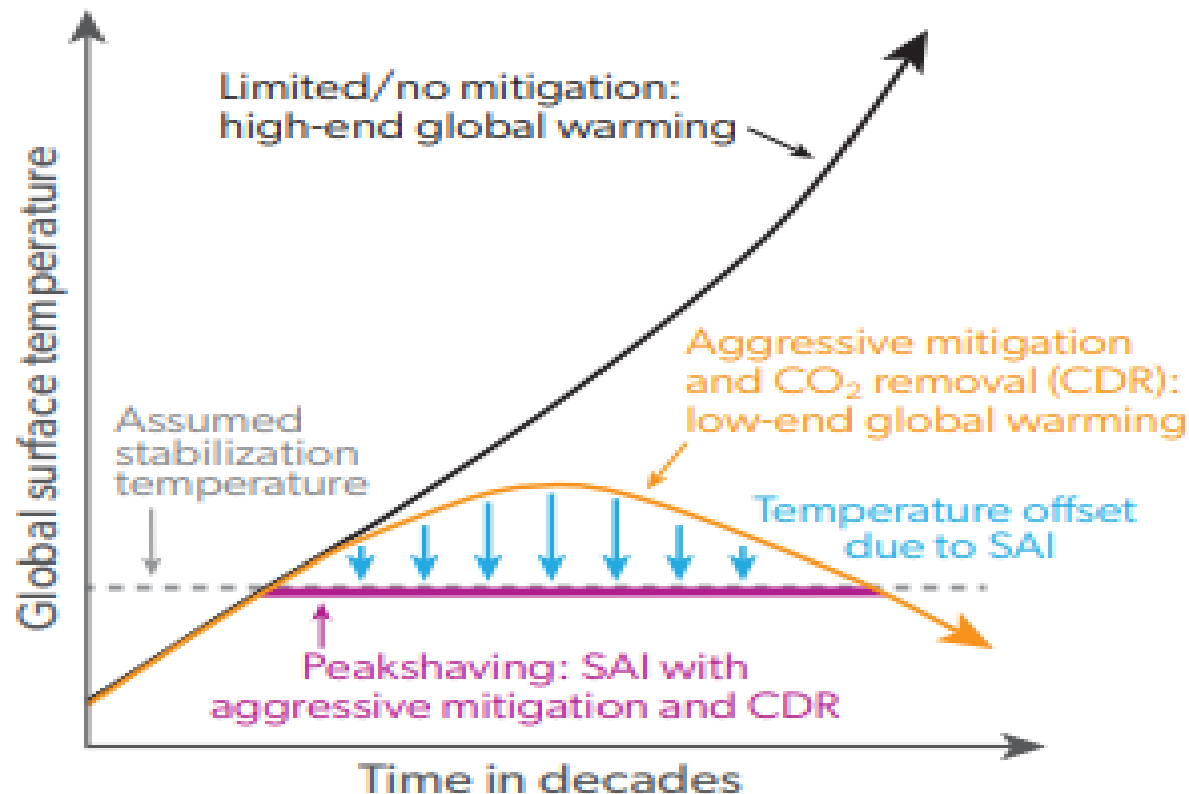
Ecosystems are not good at dealing with such shocks (Trisos et al., 2018)



Aspects such as the termination effect led to the development of ‘peak-shaving’ scenarios.

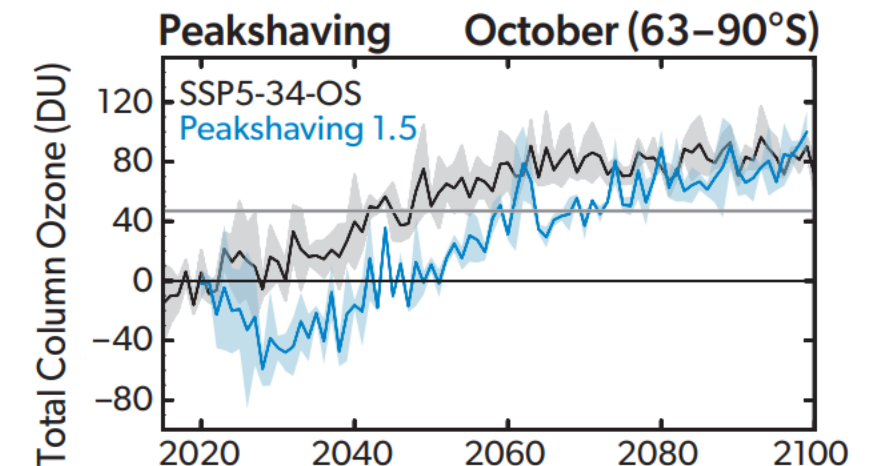
A) Peakshaving:

Aggressive mitigation and CO₂ removal (CDR) plus SAI to prevent target temperature overshoot



Current thinking is that SAI might be used in a last ditch effort together with aggressive mitigation to reduce global mean temperature over-shoot (e.g. Tilmes et al., 2020).

There are inevitably impacts on ozone: e.g. the Antarctic Ozone hole (but lots of uncertainty owing to few modelling studies).



State-of-the-science:



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Multi-model assessments reveal that we're still **very uncertain about the basics** such as i) the radiative forcing and ii) the cooling per unit SO₂ injection

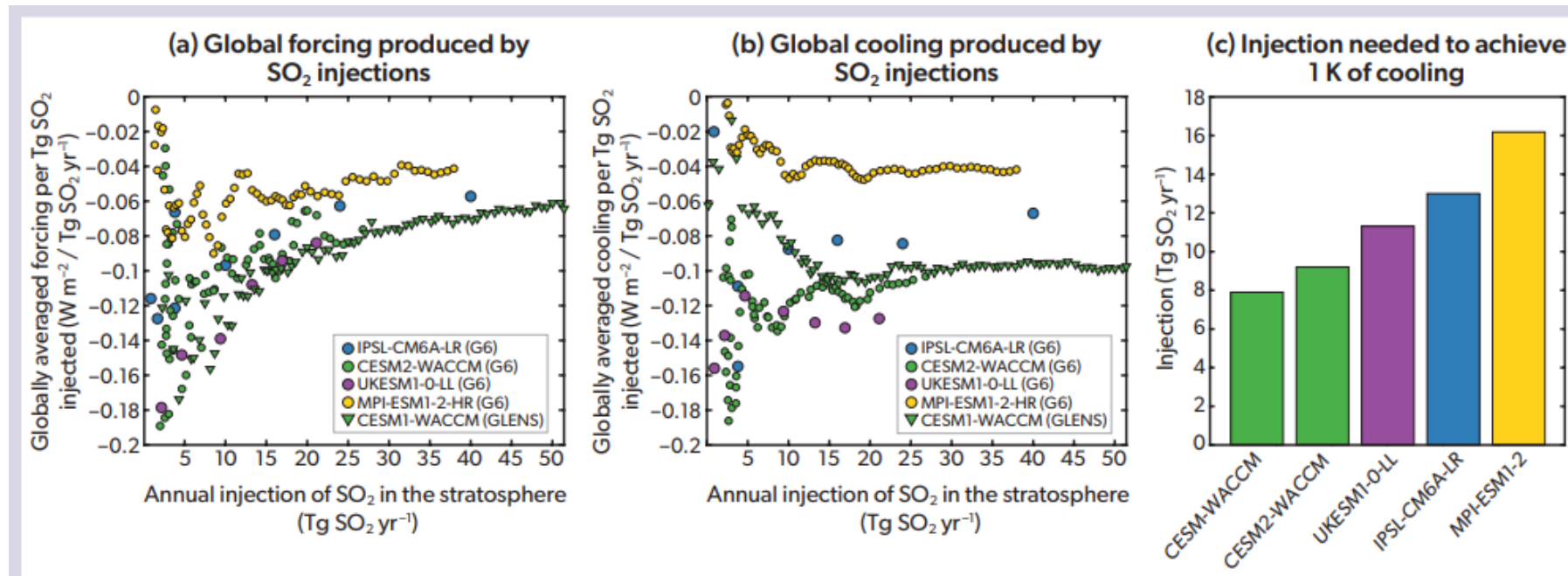


Figure 6-5. Radiation and temperature response from different Earth system model simulations of the GeoMIP scenario G6sulfur (medium SAI) and of CESM1 (WACCM) simulations in the GLENS (strong SAI) scenario. The injection rate is in Tg SO₂ yr⁻¹ (after Visioni et al., 2021a). (a) Globally averaged top-of-the-atmosphere all-sky radiative forcing (which includes the response of all ESM components) normalized by the SO₂ injection rates. (b) Annual mean global surface temperature anomaly normalized by the SO₂ injection rates. A five-year running mean has been applied to the results. (c) Injection rate of SO₂ needed to cool the globally averaged surface temperature by 1 K.

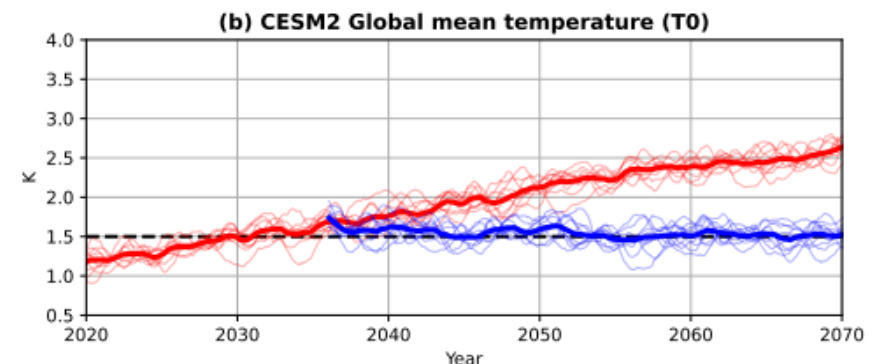
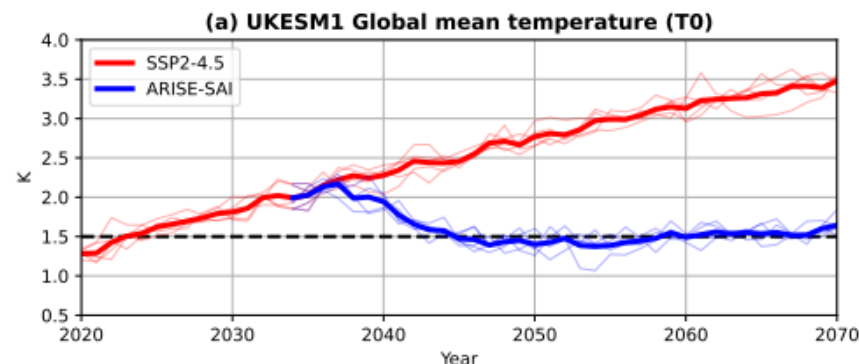


State-of-the-science:

1) Multiple injection latitudes

2) Maintaining a specified temperature target (e.g. 1.5°C, 2°C) above pre-industrial plus maintaining the hemispheric temperature gradients and the equator-pole temperature gradient

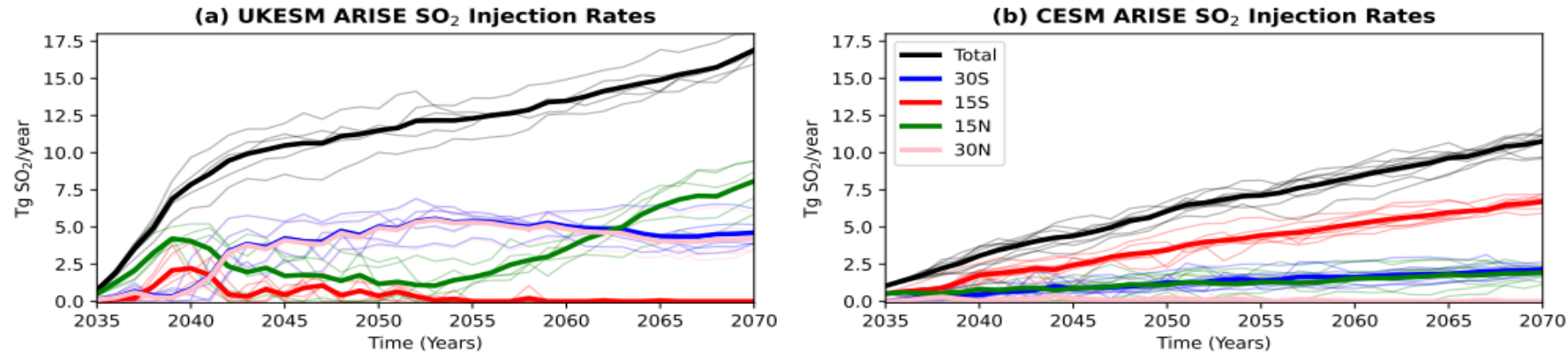
3) A controller to adjust injection latitudes as the decade progresses



Issues!

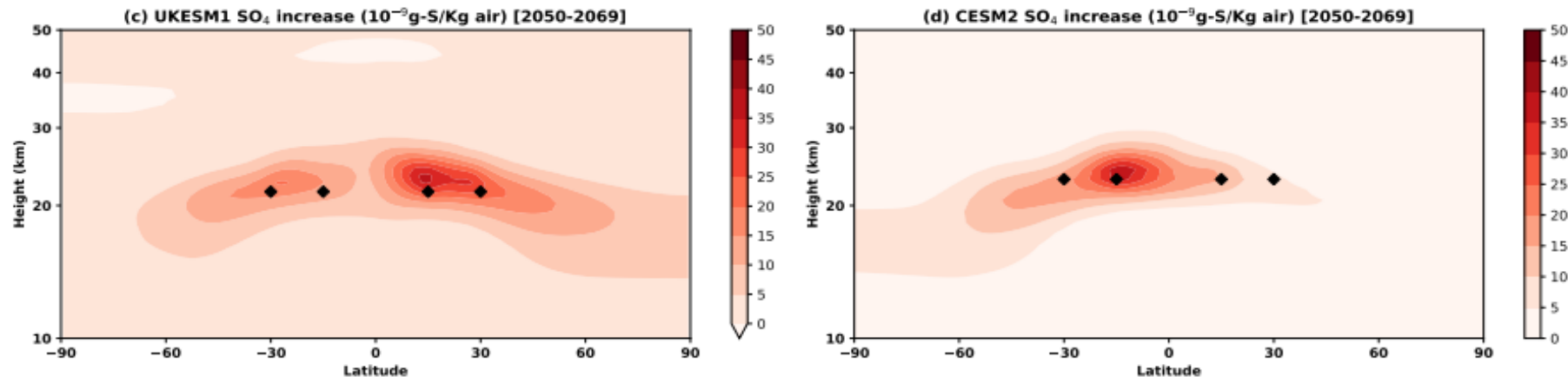


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Injection locations
wildly different!

Figure 2. Comparison of injection rates at four different latitudes in the stratosphere for UKESM1 and CESM2. Thin lines represent indi-



Aerosol optical
depths wildly
different!

More coordinated global modelling is needed to arrive at a consensus on even the basic optimal injection strategy



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MCB: The state of the science global modelling



State-of-the-science:

- Early work simply adjusted the cloud-droplet concentration (CDNC) to an observational asymptotic maximum (e.g. 375cm^{-3} e.g. Jones et al., 2009)
- GeoMIP simulations:
 - G4cdnc simply seeded increased the CDNC by 50% over all ocean areas (Stjern et al., 2018)
 - G4seasalt injected sea-salt over 30N-30S (Ahlm et al., 2017)
- New simulations
 - Explicit representation of sea-salt injection
 - Better size of aerosol injection
 - Targeting areas with high cloud susceptibility

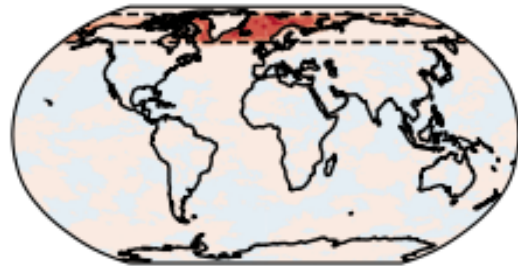
We need to understand the climate response across multiple models



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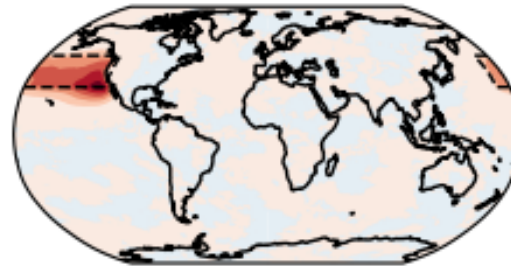
Change in the Cloud Droplet Number Concentration

Δ CDNC: 50 Tg/yr, Northern Ocean



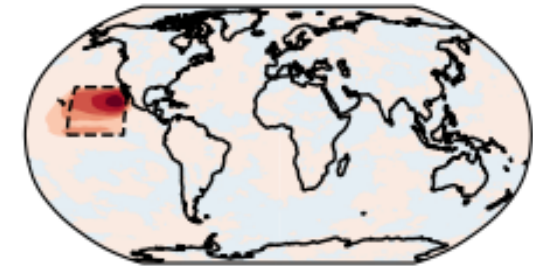
NO Δ = 247.3, abs = 311.4 cm^{-3}

Δ CDNC: 50 Tg/yr, North Pacific



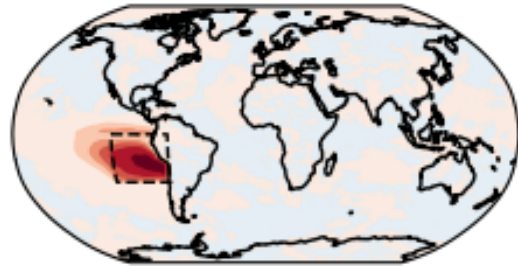
NP Δ = 315.1, abs = 401.6 cm^{-3}

Δ CDNC: 50 Tg/yr, North-East Pacific



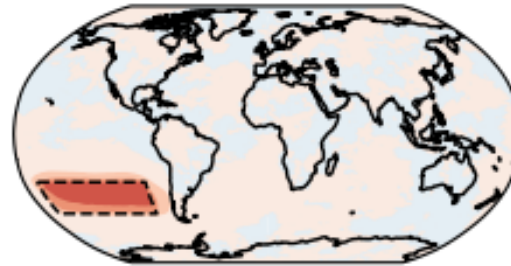
NEP Δ = 301.9, abs = 422.1 cm^{-3}

Δ CDNC: 50 Tg/yr, South-East Pacific



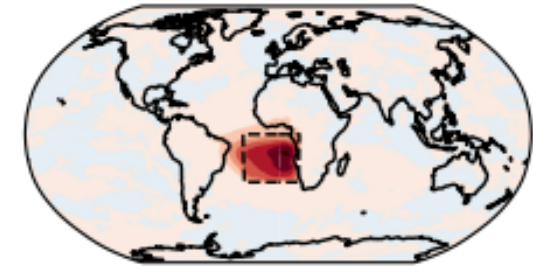
SEP Δ = 378.0, abs = 497.3 cm^{-3}

Δ CDNC: 50 Tg/yr, South Pacific

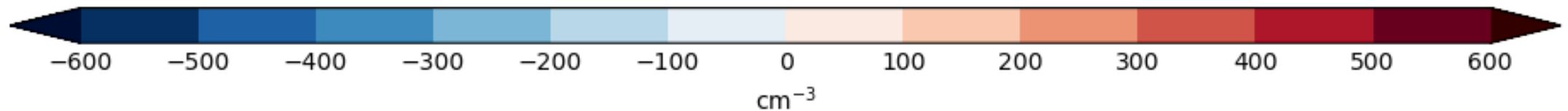


SP Δ = 304.9, abs = 358.3 cm^{-3}

Δ CDNC: 50 Tg/yr, South-East Atlantic



SEA Δ = 397.2, abs = 522.1 cm^{-3}



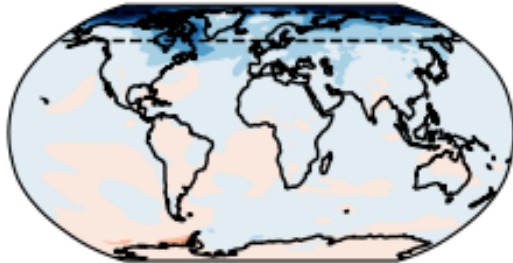
We need to understand the climate response across multiple models



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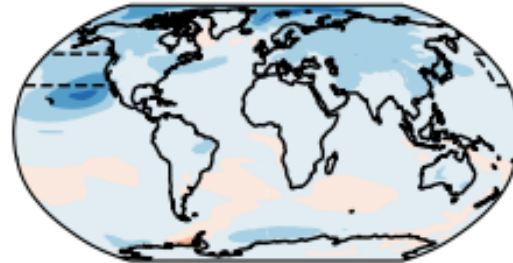
Change in the Near Surface Temperature

ΔT : 50 Tg/yr, Northern Ocean



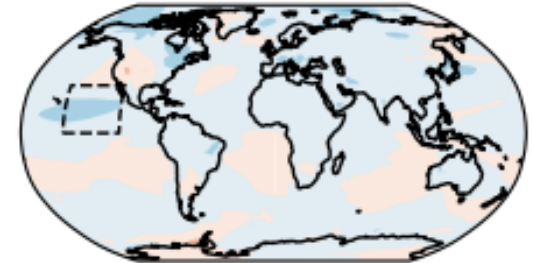
Global mean = $-0.34\text{ }^{\circ}\text{C}$

ΔT : 50 Tg/yr, North Pacific



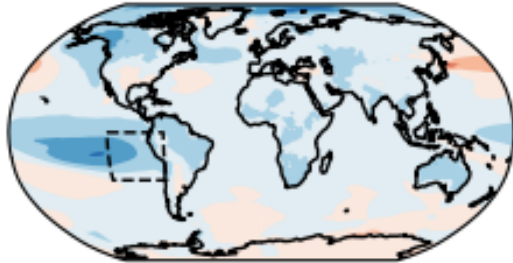
Global mean = $-0.56\text{ }^{\circ}\text{C}$

ΔT : 50 Tg/yr, North-East Pacific



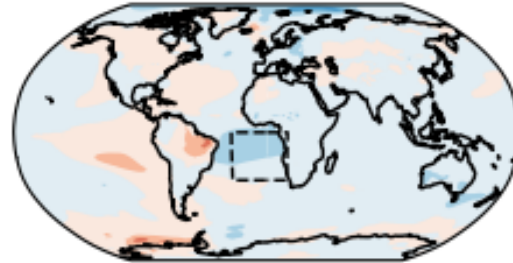
Global mean = $-0.27\text{ }^{\circ}\text{C}$

ΔT : 50 Tg/yr, South-East Pacific



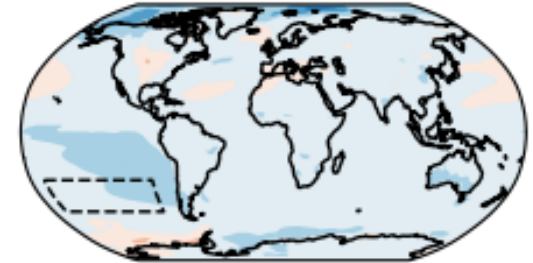
Global mean = $-0.55\text{ }^{\circ}\text{C}$

ΔT : 50 Tg/yr, South-East Atlantic

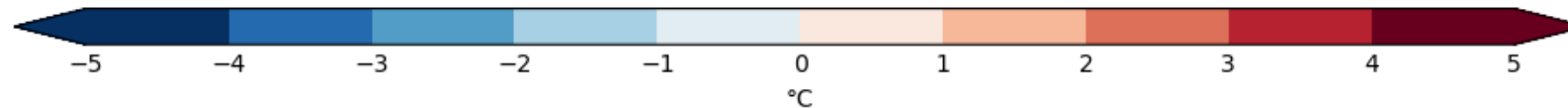


Global mean = $-0.21\text{ }^{\circ}\text{C}$

ΔT : 50 Tg/yr, South Pacific



Global mean = $-0.52\text{ }^{\circ}\text{C}$



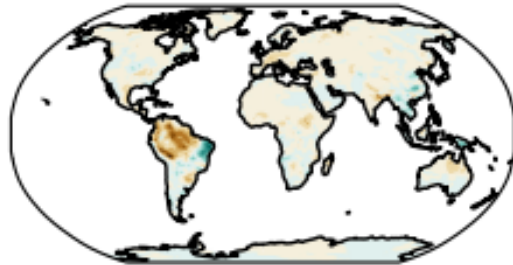
We need to understand the climate response across multiple models



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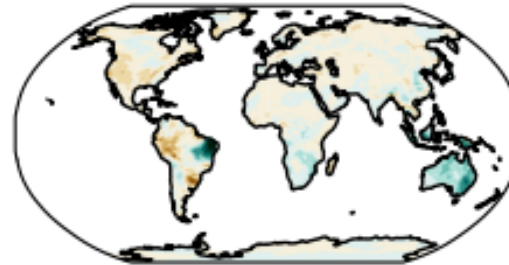
Change in the Precipitation

Δ p_{pn}: 50 Tg/yr, Northern Ocean



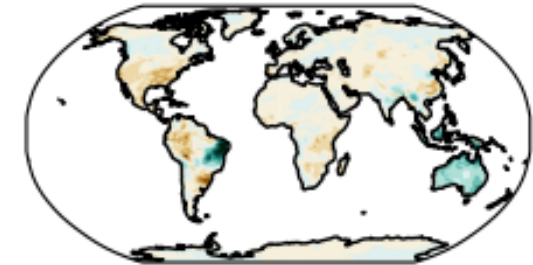
Land mean = $-0.066 \text{ mm day}^{-1}$

Δ p_{pn}: 50 Tg/yr, North Pacific



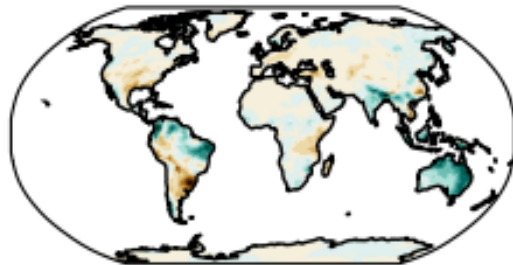
Land mean = $-0.022 \text{ mm day}^{-1}$

Δ p_{pn}: 50 Tg/yr, North-East Pacific



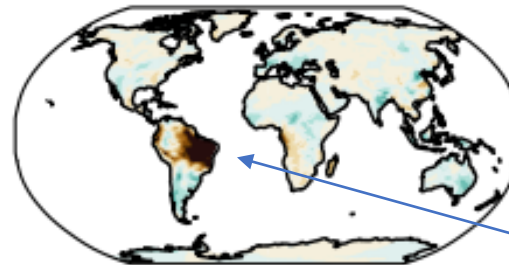
Land mean = $-0.040 \text{ mm day}^{-1}$

Δ p_{pn}: 50 Tg/yr, South-East Pacific



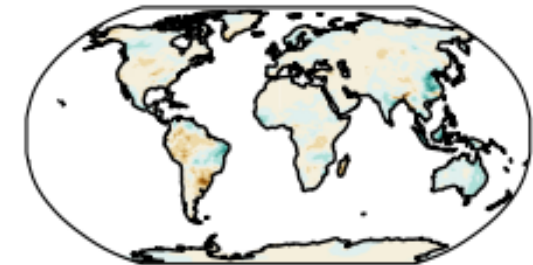
Land mean = $0.006 \text{ mm day}^{-1}$

Δ p_{pn}: 50 Tg/yr, South-East Atlantic

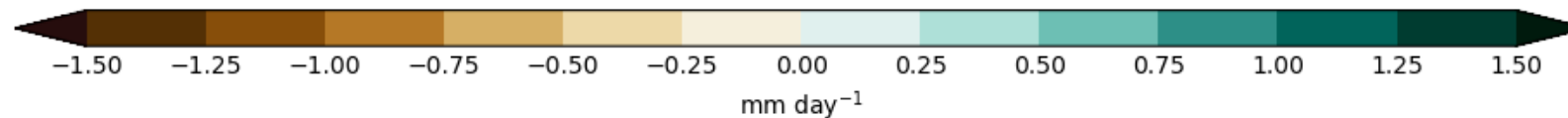


Land mean = $-0.082 \text{ mm day}^{-1}$

Δ p_{pn}: 50 Tg/yr, South Pacific



Land mean = $-0.023 \text{ mm day}^{-1}$



Does this happen
in other models?

State-of-the-science:

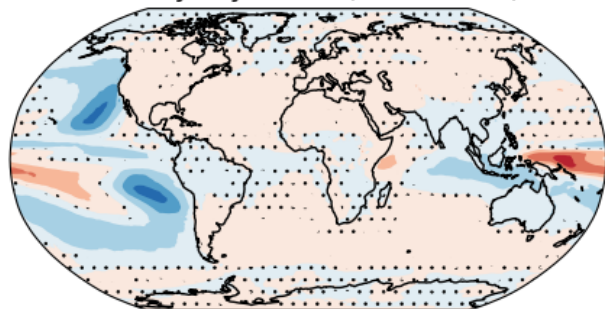


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Offsetting 1.5 W m^{-2}

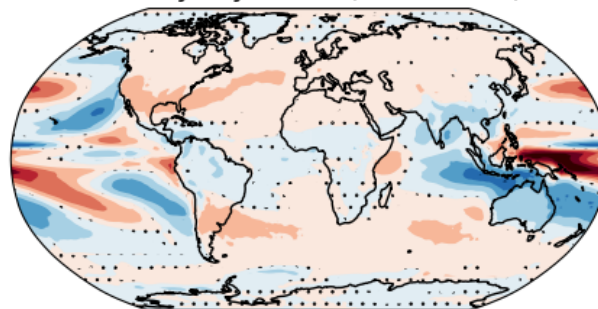
Offsetting 4 W m^{-2}

(a) Cloudy sky effect (2061-2070)



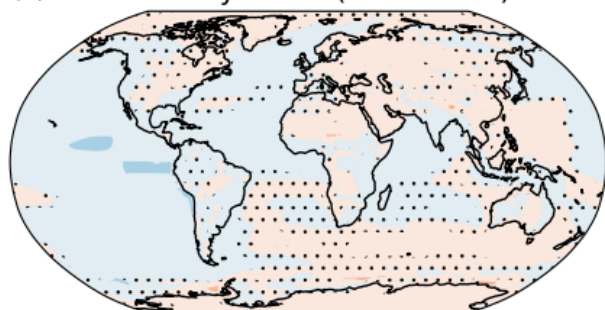
Mean = -0.80 W m^{-2}

(b) Cloudy sky effect (2091-2100)



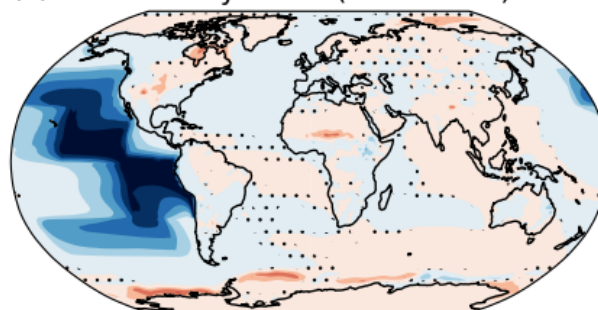
Mean = 0.32 W m^{-2}

(c) Clear sky effect (2061-2070)

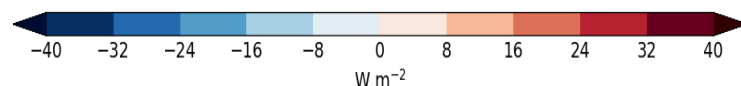


Mean = -0.67 W m^{-2}

(d) Clear sky effect (2091-2100)



Mean = -4.44 W m^{-2}



The results suggest that there could be a significant impacts from the clear-skies component – it becomes so hazy that the sea-salt itself contributes a significant amount to the brightening of the planet

BUT! Is this just an artifact of aerosol-cloud-interactions being poorly represented in global models?

Surely you need better resolution than the 100km resolution of GCMs to adequately represent cloud-scales?

Again, more coordinated global modelling is needed; but also modelling at smaller scales and better observations



Conclusions

- Brightening the planet by SAI or MCB may provide the only way to maintain global mean temperatures at e.g. 2° C above pre-industrial levels.
- Climate modelling of stratospheric aerosol injection suggest the cooling efficiency is poorly constrained and strategies to maintain global mean temperatures and ameliorate side-effects reveals radically different injection strategies.
- Climate modelling of marine cloud brightening remains extremely uncertain. There is so much complexity that is poorly represented by climate models.
- Coordinated research is desperately needed given the urgency of the climate situation.

Support Slides

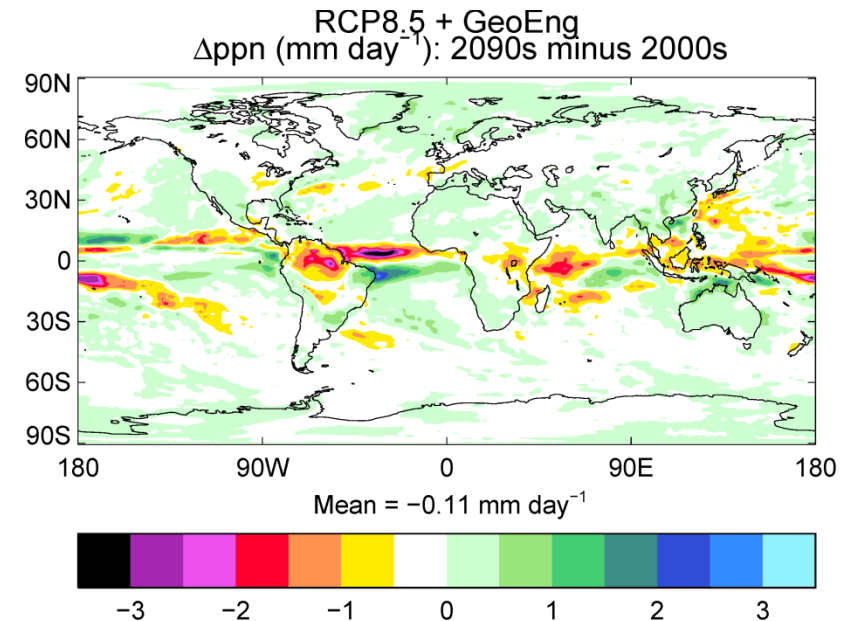
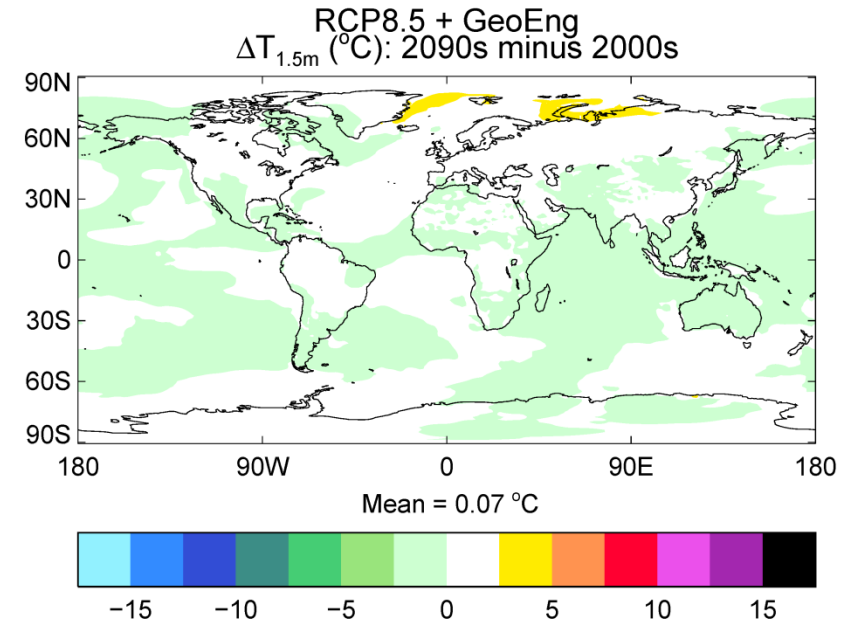
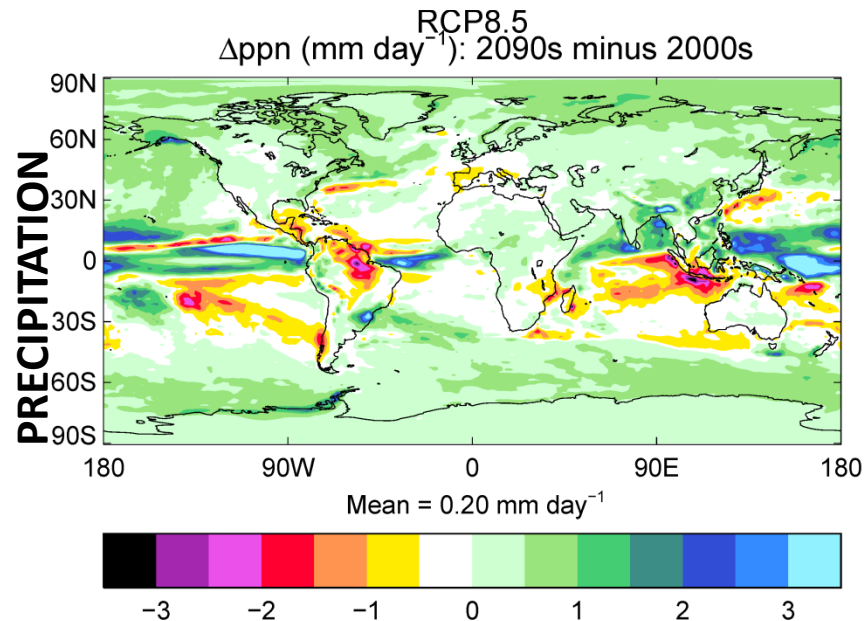
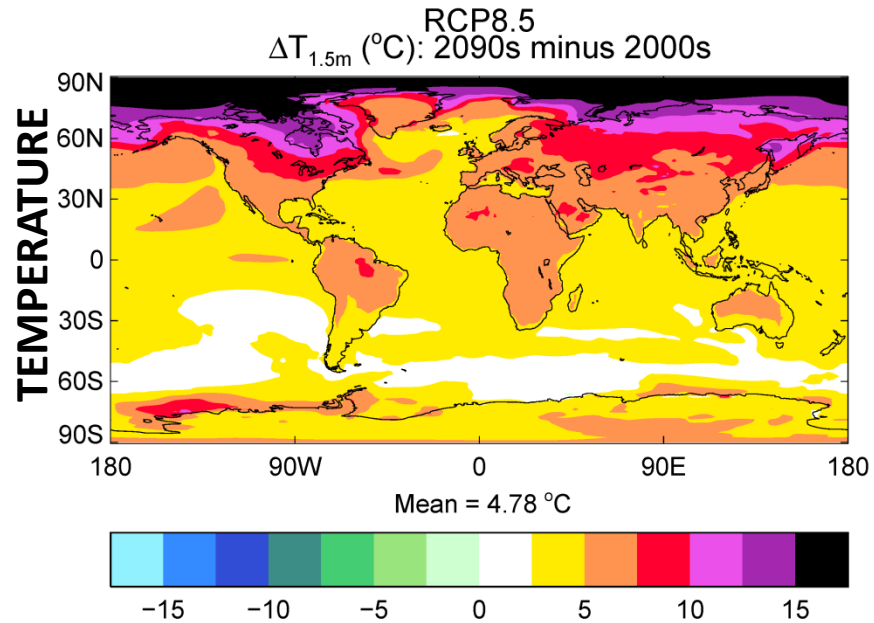


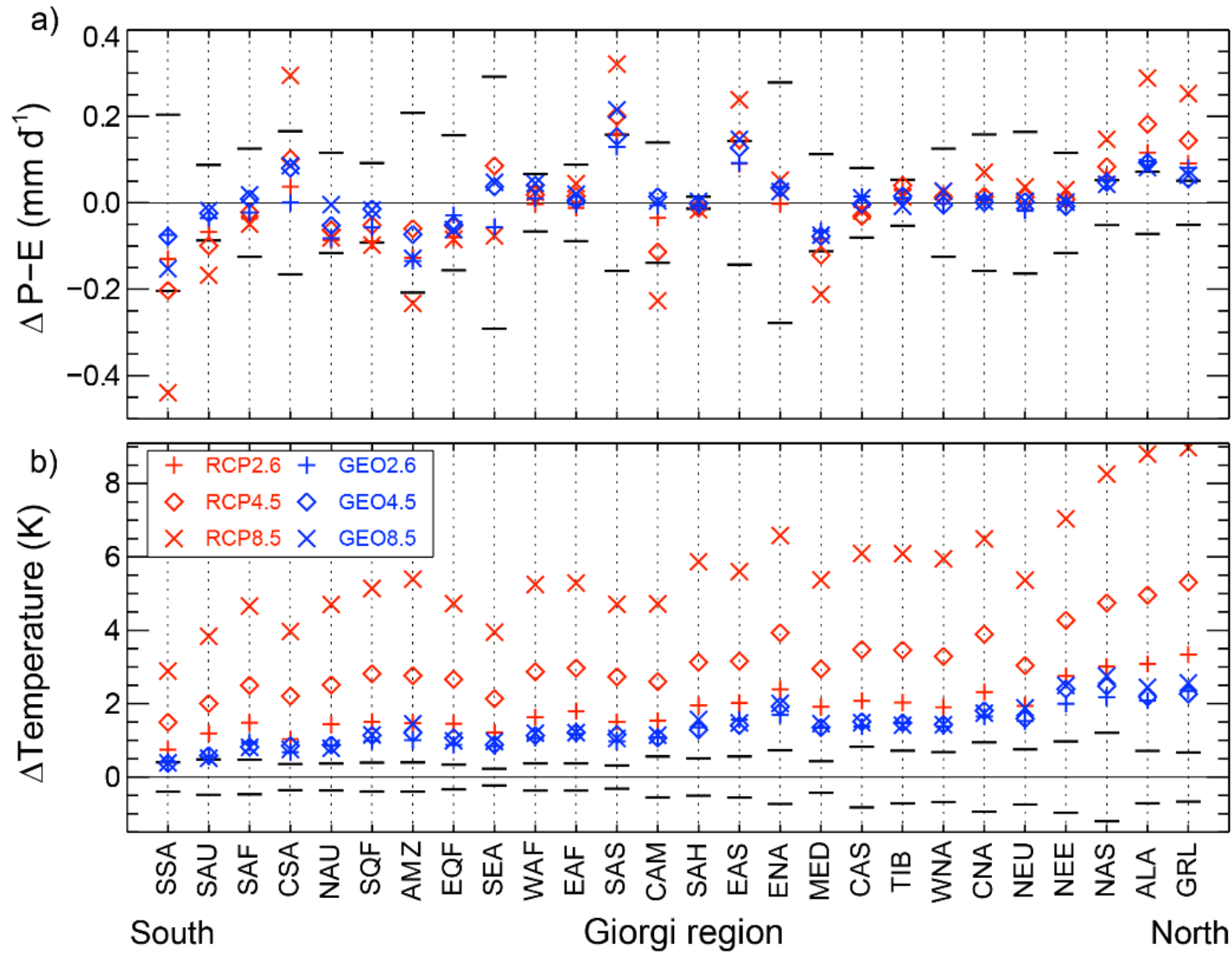
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Solar Radiation Management?

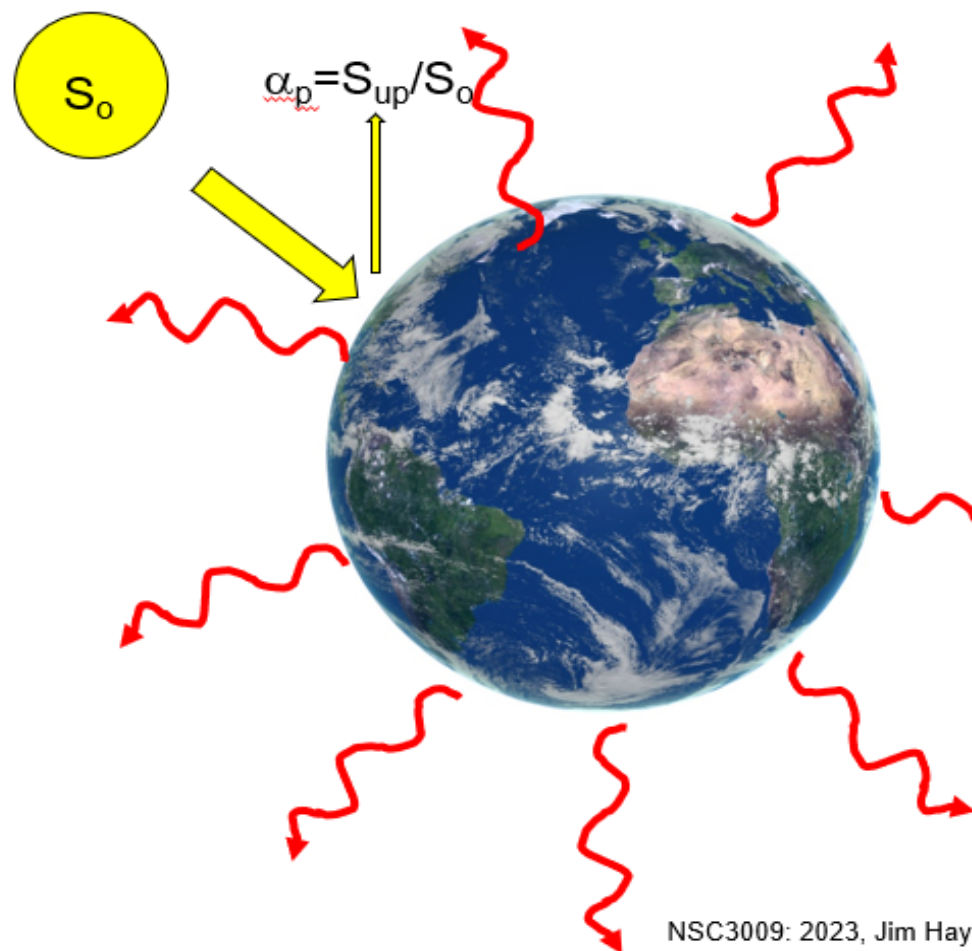


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The Earth's temperature is controlled by the balance between the sunlight absorbed by the Earth and the emitted thermal radiation



For the Earth system where, r =radius of Earth, solar constant = S_0 , α_p =planetary albedo, ε = effective emissivity:

$$\text{Sunlight absorbed} = S_0 \pi r^2 (1 - \alpha_p)$$

$$\text{Emitted thermal radiation} = 4 \pi r^2 \varepsilon \sigma T^4$$

$$\text{For balance: } S_0 \pi r^2 (1 - \alpha_p) = 4 \pi r^2 \varepsilon \sigma T^4$$

Therefore:

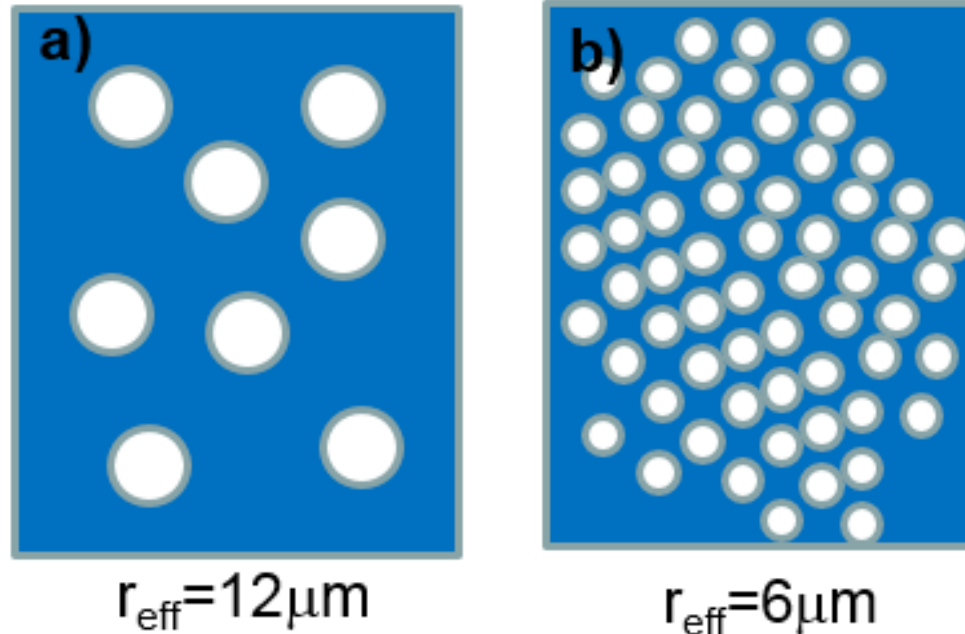
$$T = \sqrt[4]{\frac{S_0 (1 - \alpha_p)}{4 \varepsilon \sigma}}$$

So you can cool the Earth by either reducing S_0 or increasing α_p (or ε)

Indirect effects 'aerosol-cloud-interactions'

Imagine you are a satellite instrument looking down on clouds on Earth.

Suppose both have the same liquid water content (volume proportional to r^3), but droplets in b) are half the diameter of those in a). You get 8 times more cloud droplets. Hence b) is more reflective than a)



MCB will work most effectively if there is a reduction in r_{eff} AND an increase in LWP AND an increase in cloud fraction/cloud lifetime