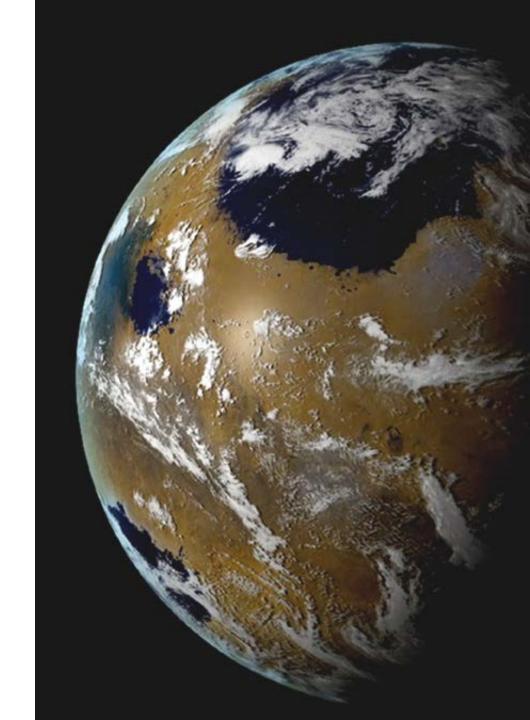
The Early Mars Climates Enigmata

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Sorry: no references are cited in this presentation (except figures)



Which atmospheres for early Mars?

- Direct constraints from rocks
 - From rock samples (meteorites, MSR). Gas inclusion?
 - Clues from geology and mineralogy
- Constraints from our understanding of the formation of the Noachian atmosphere
 - Most likely no atmosphere accumulated before 4.1 Ga because of early EUV heating (possible sporadic atmosphere sduring strong volcanic degassing)
 - ⇒ To be confirmed by geo observations?
 - After 4.1 Ga: which atmosphere could form from late outgassing & impacts?
- Constraints from our understanding of the loss of the Noachian atmosphere
 - Escape
 - Crustal reservoir sinks (not only carbonates?)

Global mean Temperature

Modeling the Noachian-Hesperian ClimateS

- 1. 1D global radiative convective models
- ⇒ To evaluate global mean surface temperature with various atmospheres
- 2. 3D General Circulation model
- To evaluate local surface temperatures & their variations with season, obliquity, the role of clouds, atmospheric circulation etc...
- 3. 3D Global Climate model with realistic surface reservoirs produced by a converged water cycle with hydrology/cryology models
- ⇒ Necessary to evaluate rain, snow melting, runoff, lakes evolution, etc.

WARNING: Getting warm temperatures without modeling water can be misleading:

- Water affects temperatures (latent heat, thermal inertia, albedo)
- Water tends to accumulate in the coldest regions (cold traps).
 - ⇒ With a limited water inventory, a very warm climate is necessary to get liquid water in cold traps (Dune planet)

Modeling the Noachian-Hesperian ClimateS with CO₂

Modeling with a thick CO2 + N2 + H₂O atmosphere

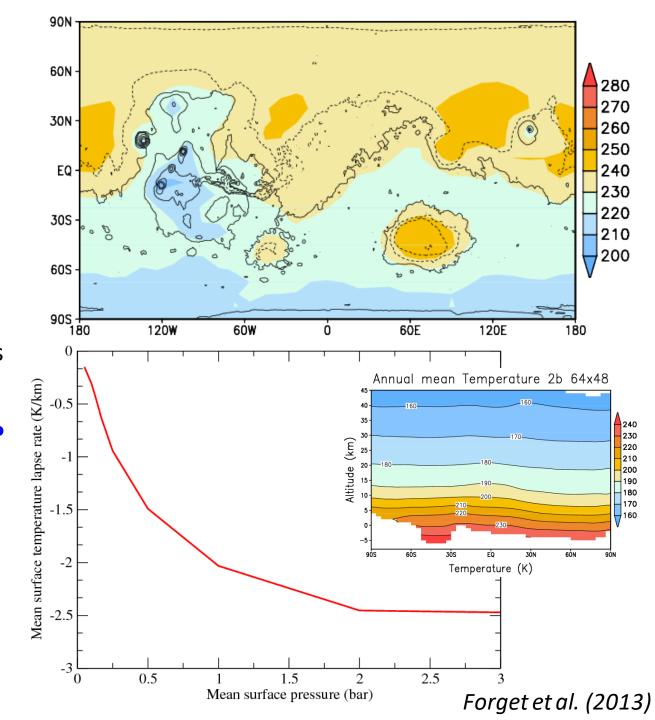
- Mars not warm enough for rivers
- With a thicker atmosphere, cold early Mars is different than today: cold highland and warm plains ⇒ "Icy highland scenarios"
- A "surprise" warmer solution could still be discovered (e.g. with improved models of clouds or aerosols) but it is unlikely.

Solution with an early Sun warmer than expected?

 not yet completely rejected (a more massive Sun is predicted by some recent theories and young star observations, but for 100 Myr, not 1Gyr)

⇒ This could be checked using

- Remaining signatures of the early solar wind
- The effects of a solar mass-loss history on the orbital dynamics in the planetary system, (signatures in the geological record?)



Warming early Mars with other greenhouse gases

- NH₃: photochemically unstable
- Volcanic SO₂ and H₂S : climate affected by aerosols
- Exotic ideas: N₂O, H₂O₂, CF3Br, CF2Cl2, ...
- The new trend: H₂ & CH₄?
 - Need a background pressure (greenhouse effect only induced by collision induced absorption or "CIA") >> 0.5 bars
 - Recent progress in the H₂ & CH₄ CIA absorptions coefficients measured in the lab.
 - H₂ escapes quickly, CH₄ is photolized ⇒ Significant sources are required :
 - Highly Reduced mantle ? ⇒ Outgassing of H2 directly and/or CH4, later photolized in H₂. Then what about CO₂?
 - Serpentinization?
 - Production of H₂ and CH₄ in the thermal plumes of impacts ?
 - ⇒ What evidences are in the rock records?

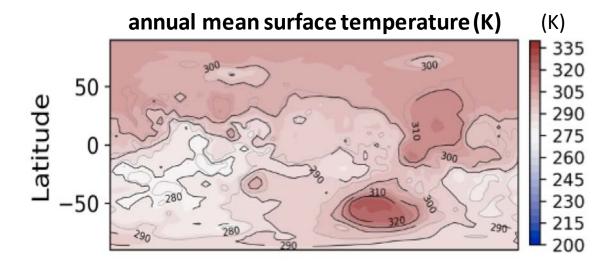
Modeling the Noachian-Hesperian ClimateS with CO₂ + H₂

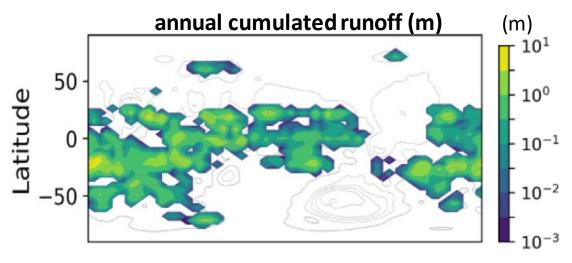
 Recent 3D simulations by Global Climate Models

Ames, ROCKE-3D, LMD, PMGCM (Japan)

⇒promising results, with possible runoff over valley network locations for fundamental reasons.

- Large parameters space to explore (obliquity, orbit, water inventory, atmospheric composition, topography, model parameters, etc.)
- Models are not yet complete enough to simulate the evolution of lakes, glaciers, the formation of seas, etc...





The climatic impacts of impacts

- A few large ancient impacts (crater >> 100 km)
 - should induce short (tens of years) warming, with deluge-style precipitation (unless H₂ CH₄ are produced?)
 - can affect the geo heat flux for millions of years
 - happened earlier than valley networks
 - BUT: strong erosion and alteration, regolith layers formation
- Many more smaller impacts (10 km < craters < 100 km)
 - production of ice-melting ejecta layers ?
 - maximum melt production ~10-1000 times lower than what is needed for the valley network
- YET: Large impacts must have left climate-induced signatures in the geological records.
- Could impacts explain very late fluvial forming events?



Inputs needed from geological studies

- Modelers need more constraints from geomorphology and mineralogy
 - sometime observations not useful by themselves can constrain climate models
- <u>Consensual</u>, quantitative (corrected from observational biases) measures are useful (even with debates on the interpretation).

Examples:

- Reliable map of terrains with a given age with or without fluvial features.
- Detailed temporal relationship between fluvial forming events/ craters / volcanoes
- Presence of glaciers on Noachian & Hesperian Mars ?
- Runoff estimation, open-closed-basins distributions, etc.
- We must elucidate the links between hydrous minerals and climate, and their timing
 - Example: Which philosilicate are produced by surface liquid water, hydrothermalism, impacts,
 Primordial alteration by a steam atmosphere, etc.
- What was Mars topography through time (True Polar Wander ?)

Final comment

Using terrestrial analogs and experience can be misleading on Mars

- On recent Mars: wrong interpretation of gullies, RSL, spiders
- On early Mars: exotic process must be imagined and further studied:
 - Examples:
 - Fast erosion in impacts-made regolith
 - Signature of intense precipitation after impacts
 - Geomorphology in a world were eolian erosion dominates fluvial erosion for extremely long duration...
 - Signature of CO₂ glaciers...
- ⇒ Numerical modeling and lab experiments can help