

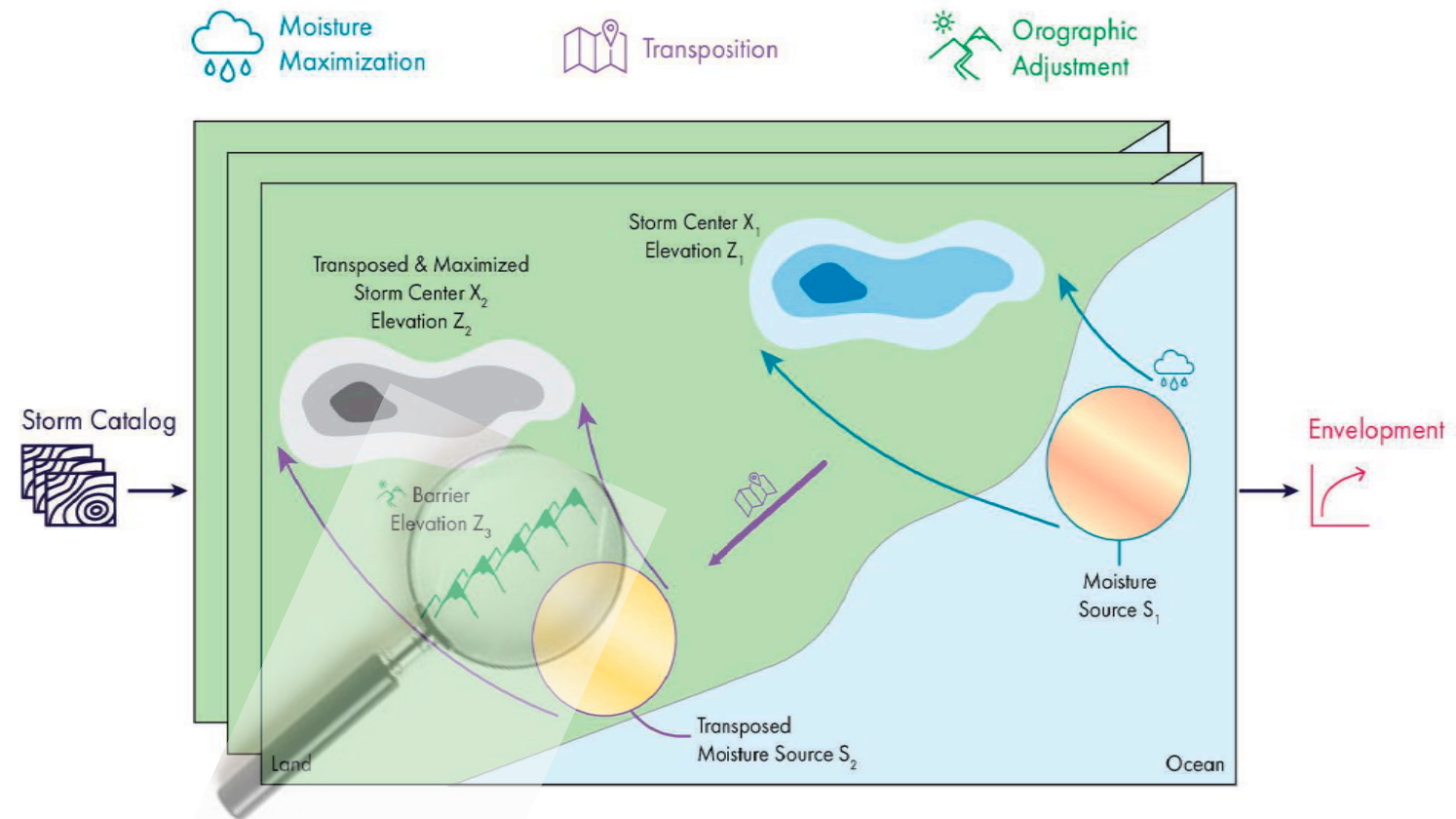
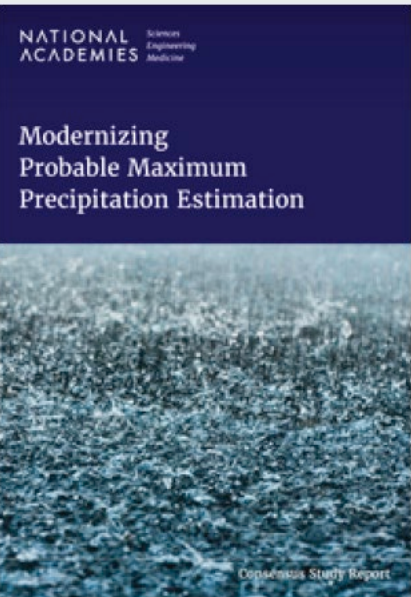
Science Challenges for Modeling Extreme Warm Season Rainfall in Mountainous Terrain

Richard Rotunno



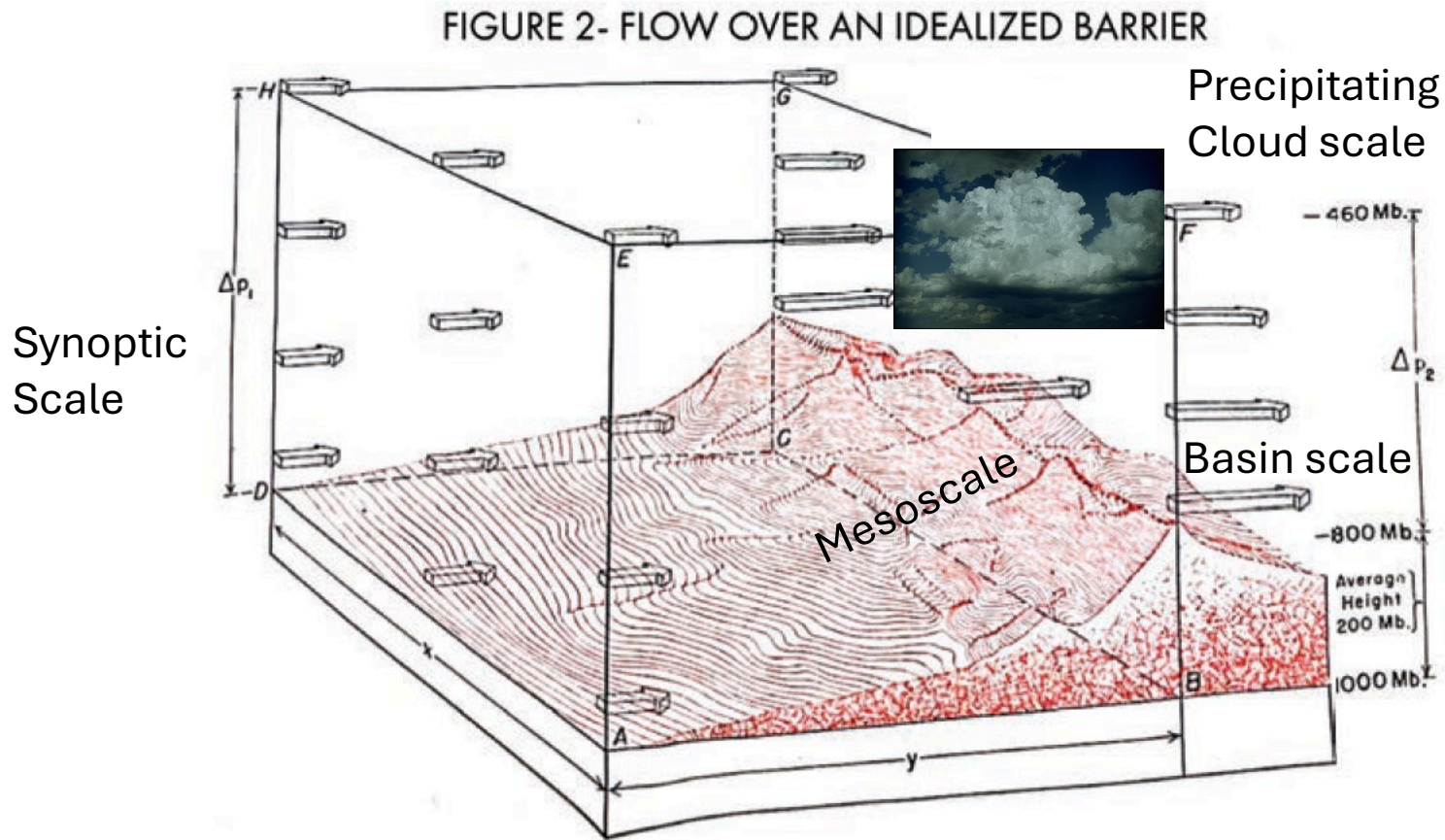
Probable Maximum Precipitation

2024



Probable Maximum Precipitation

The **depth of precipitation** for a particular **duration, location, and areal extent**, such as a drainage basin, with an extremely low annual probability of being exceeded, for a specified climate period.



Depth of Precip. =
Duration x Rain rate
(Synoptic) (Clouds → Precip)

Location and Areal extent →

“It matters where it lands”

(Cloud growth and movement)

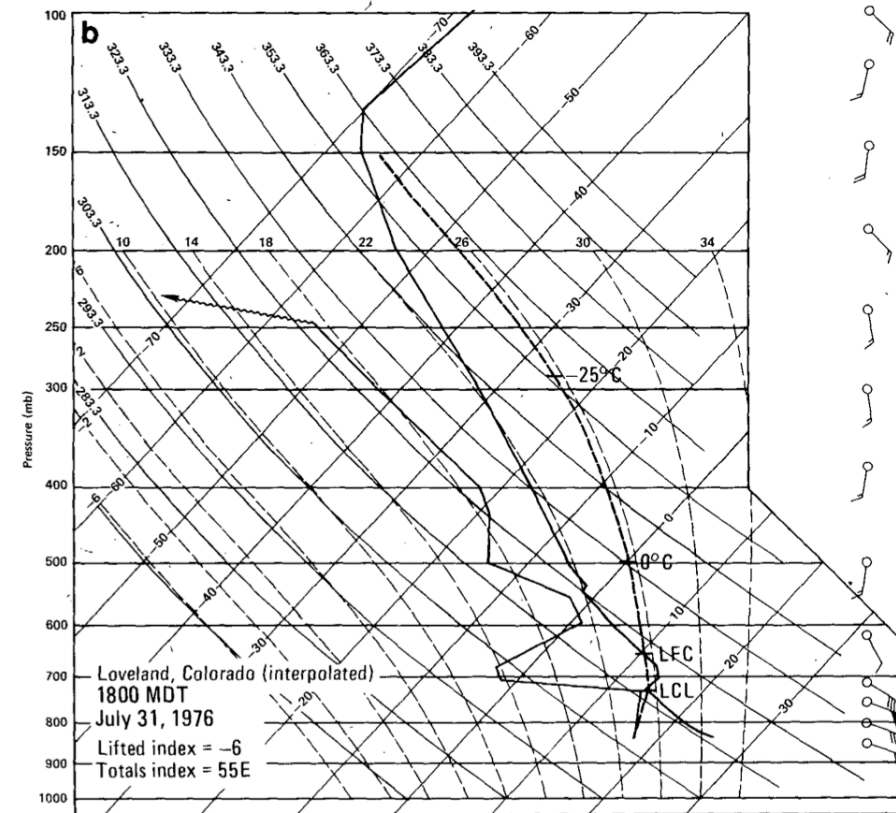
FIGURE B-2 Conceptual orographic model for PMP based flow over a ridge.

SOURCE: HMR 21B (USWB, 1945), Figures 1 and 2.

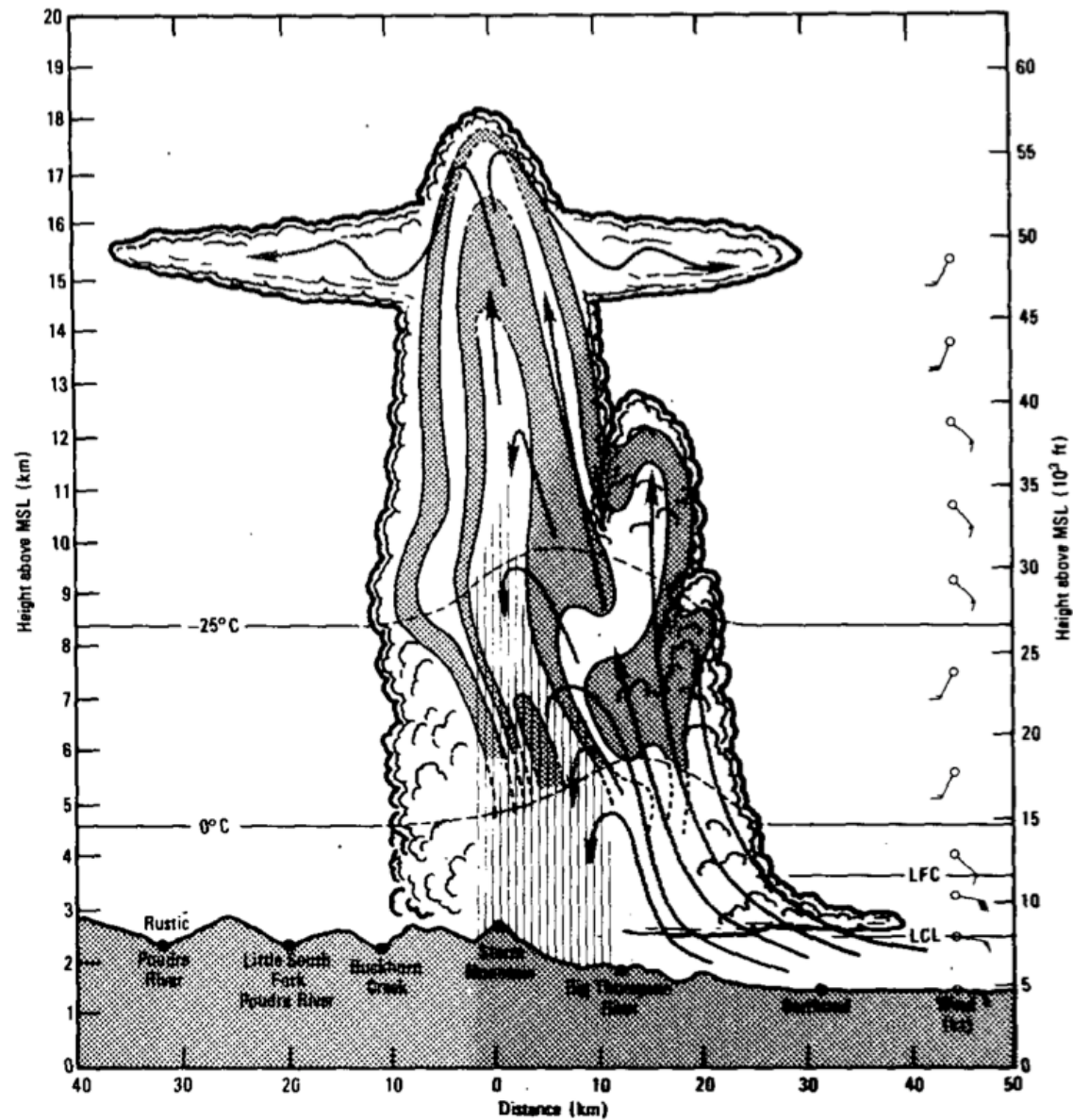
Big Thompson Flood Colorado, 1976



low-level
flow



Big Thompson Flood Colorado, 1976



Caracena et al. (1979, *Mon Wea Rev*)



170mm in several hours

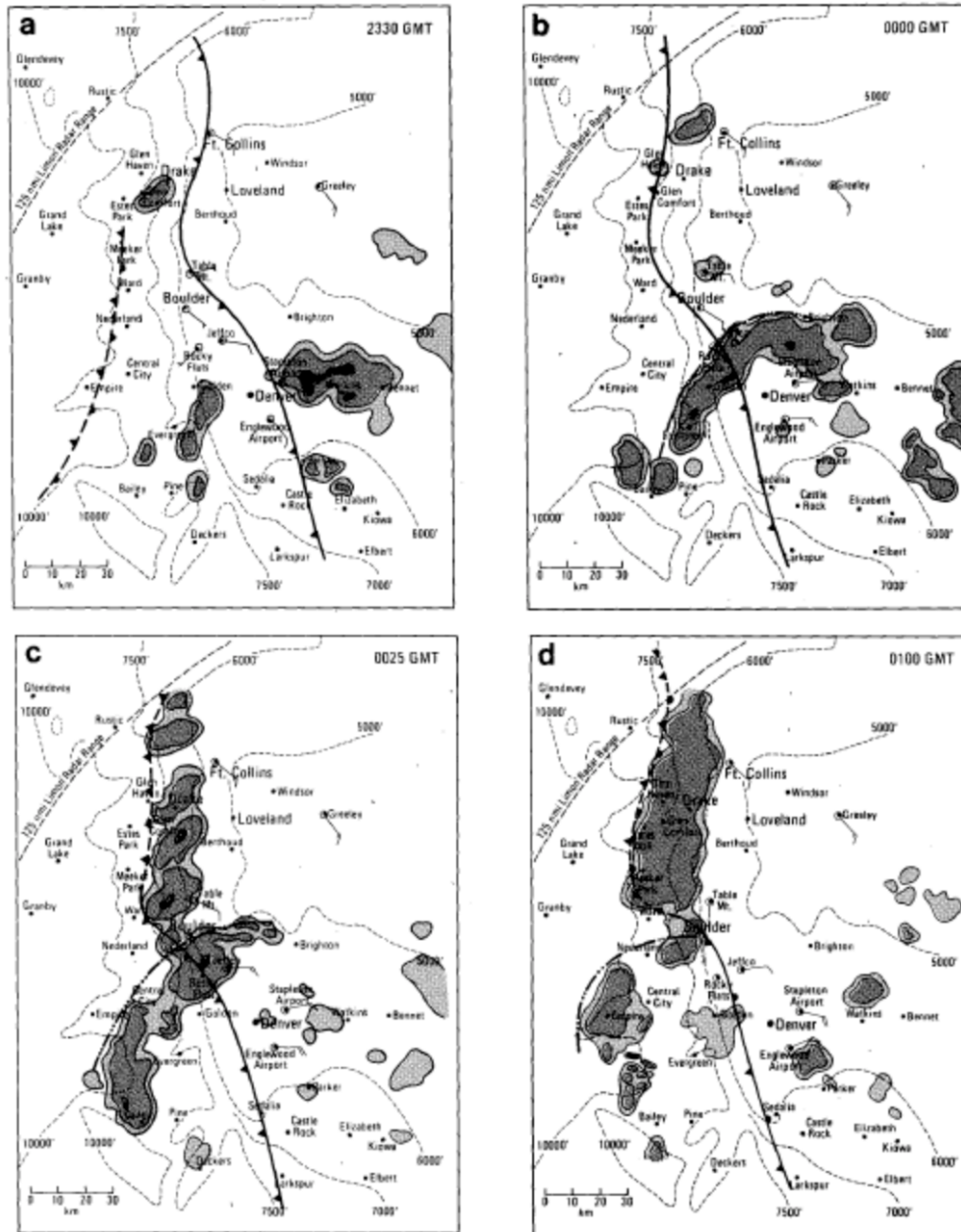
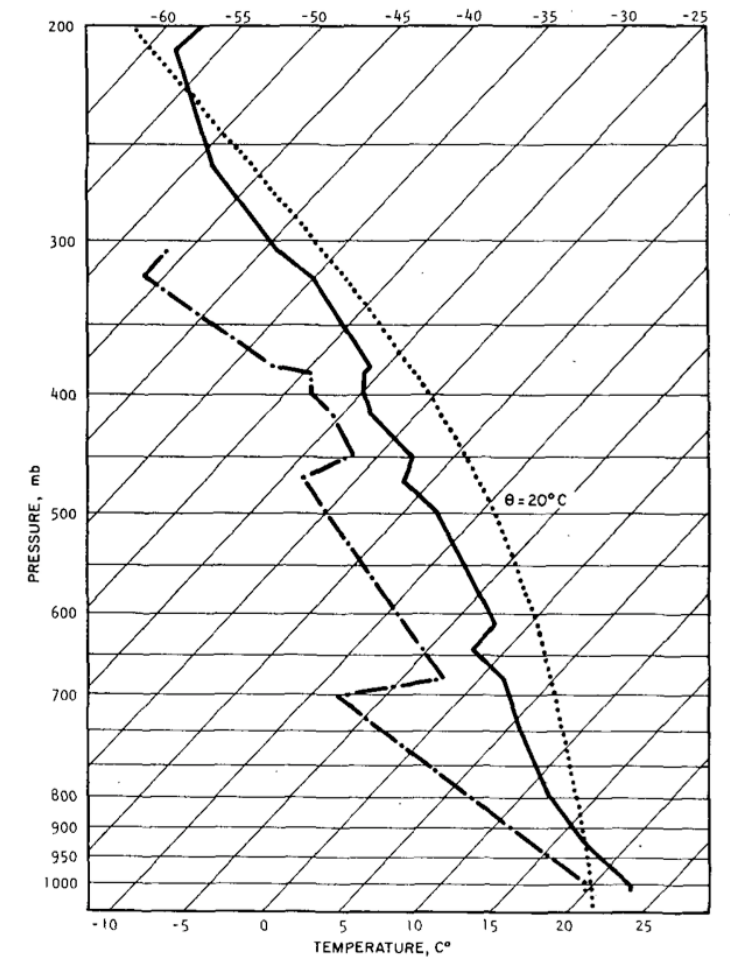
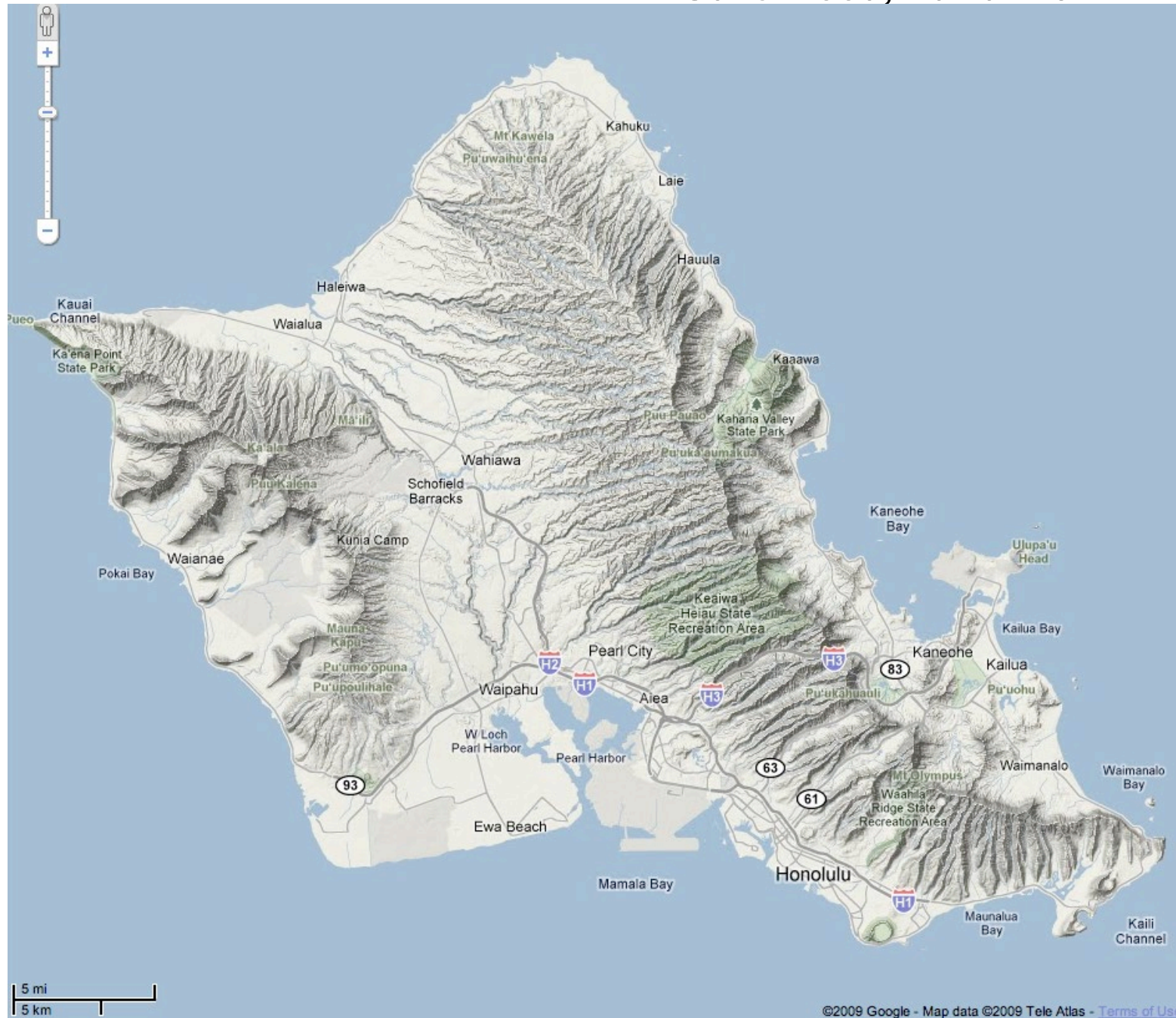


FIG. 1. Local-scale surface analyses. Frontal positions and reported winds are in black; Limon radar echoes are shown with VIP level 1 return shaded light gray, level 2 shaded medium gray and level 3 shaded dark gray. (a) 2330 GMT 31 July 1976, (b) 0000 GMT 1 August 1976, (c) 0025 GMT 1 August 1976, (d) 0100 GMT 1 August 1976.

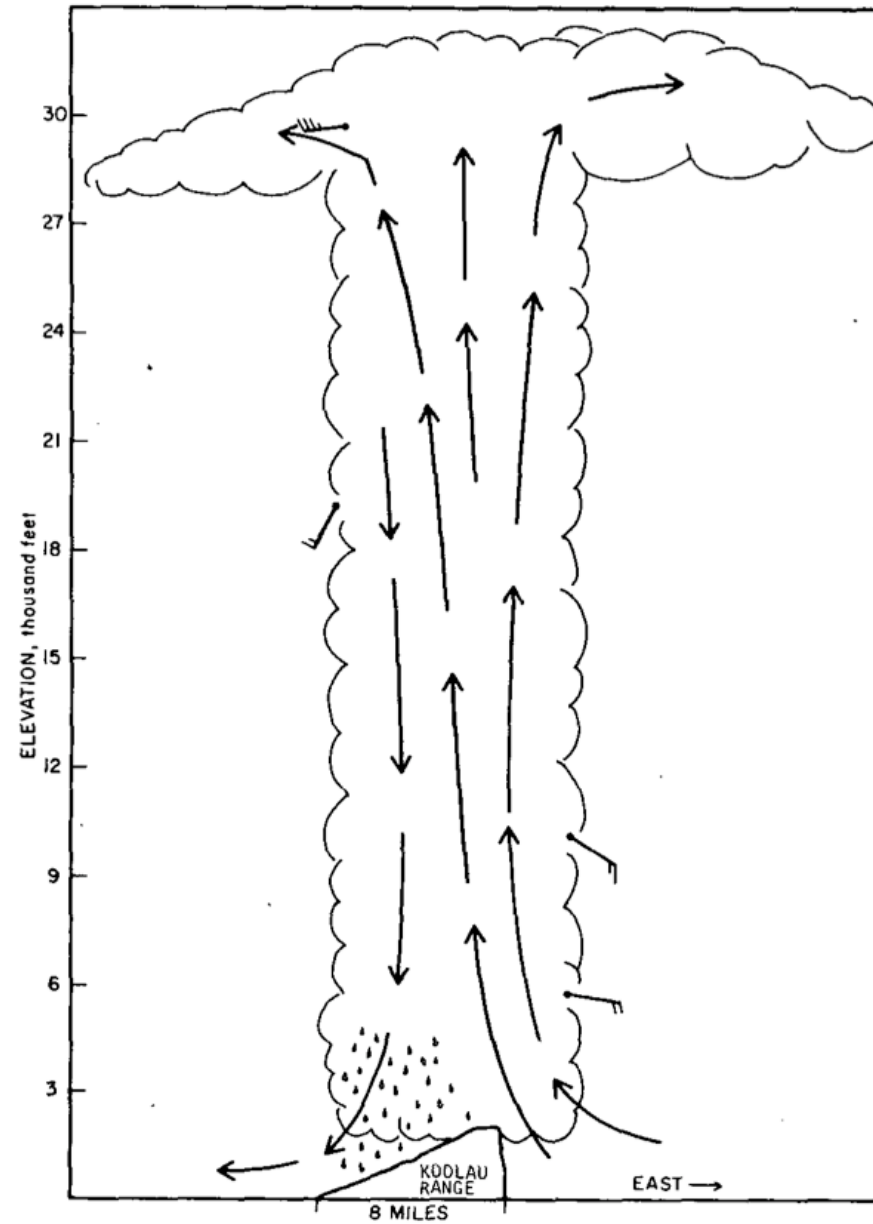
Cells drifting up from the south and amplifying over the Big Thompson River Drainage

Oahu Flood, Hawaii 1974



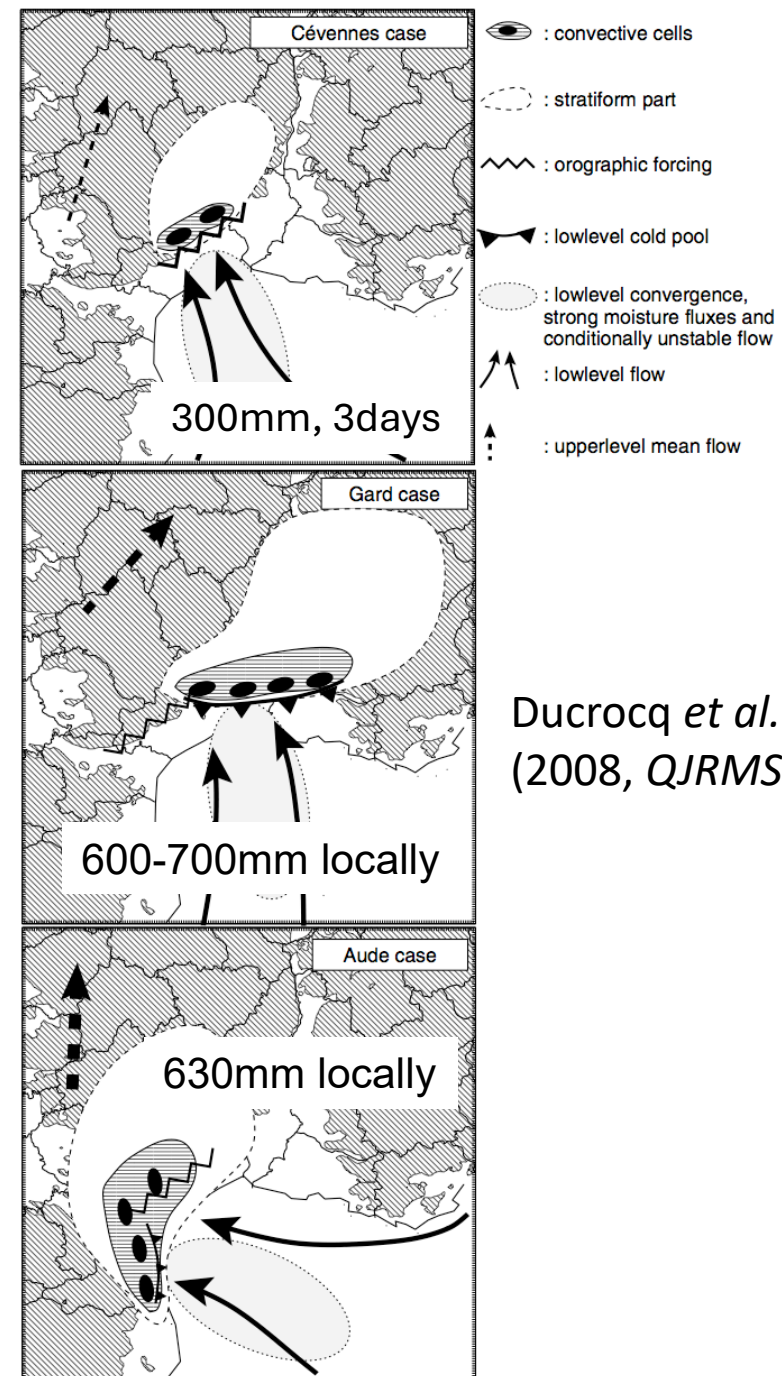
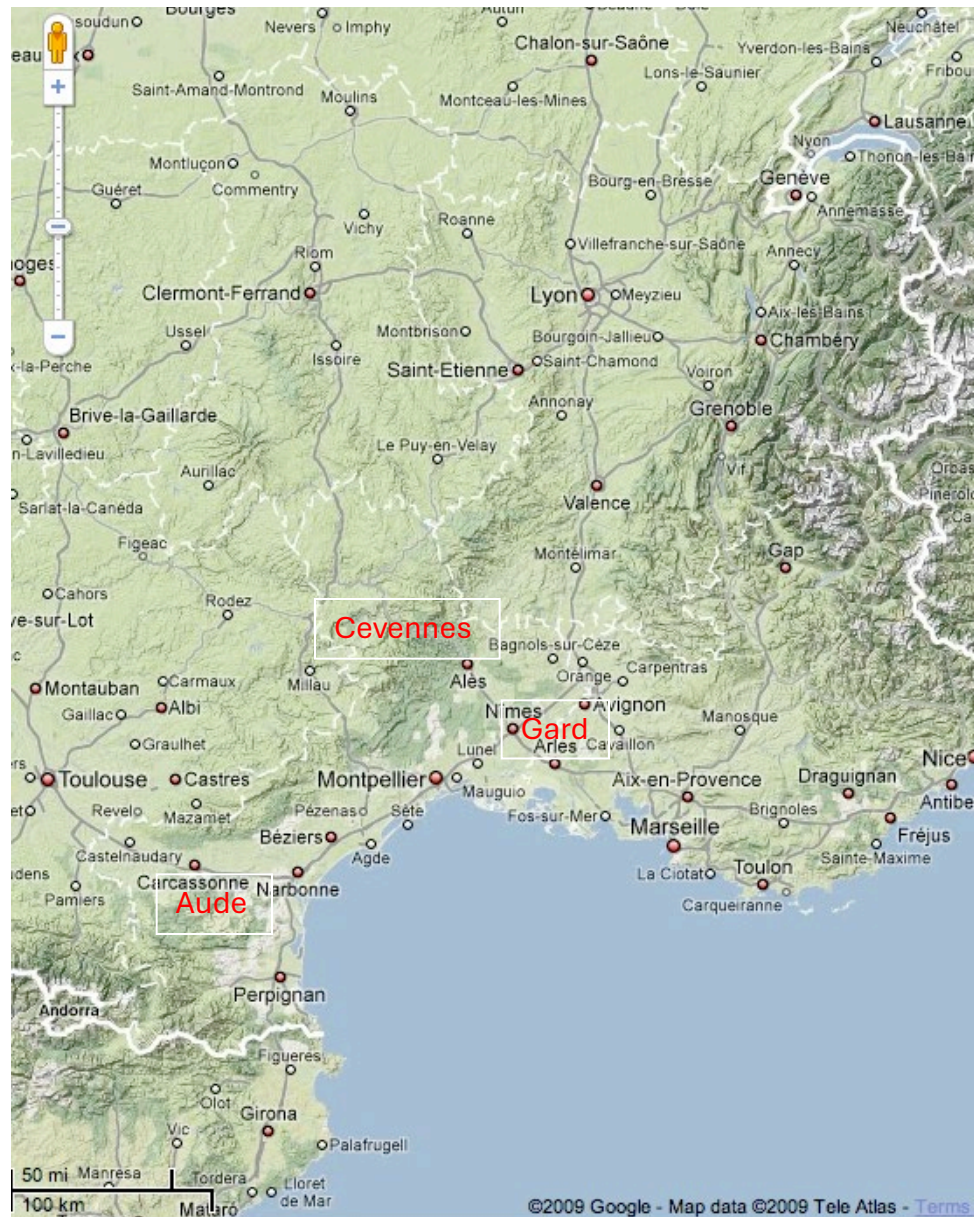
←
low-level
flow

Oahu Flood, Hawaii 1974



Schroeder (1977, *Mon Wea Rev*) 250mm in several hours

Floods in Southeastern France



Ducrocq *et al.*
(2008, *QJRM*S)

Three nondimensional parameters emerge as most important:

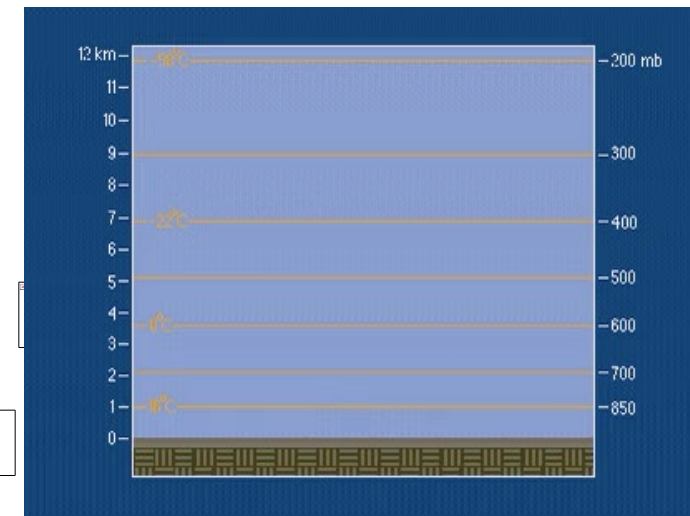
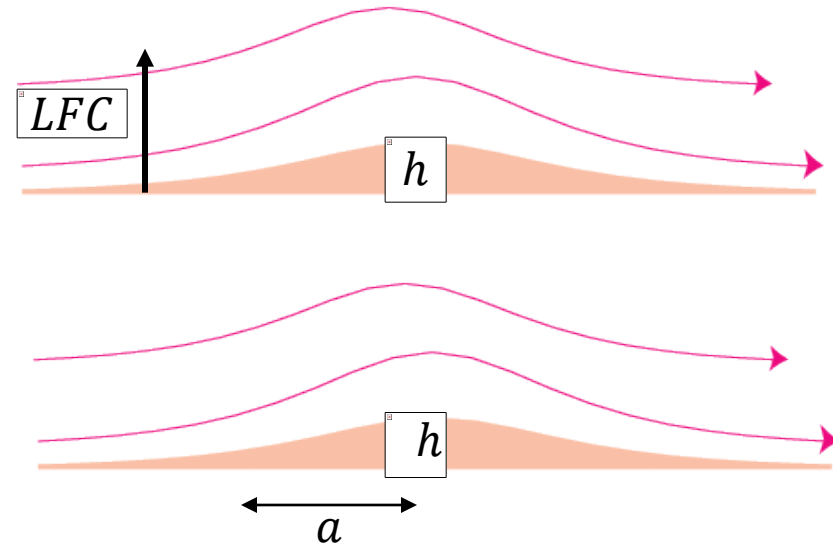
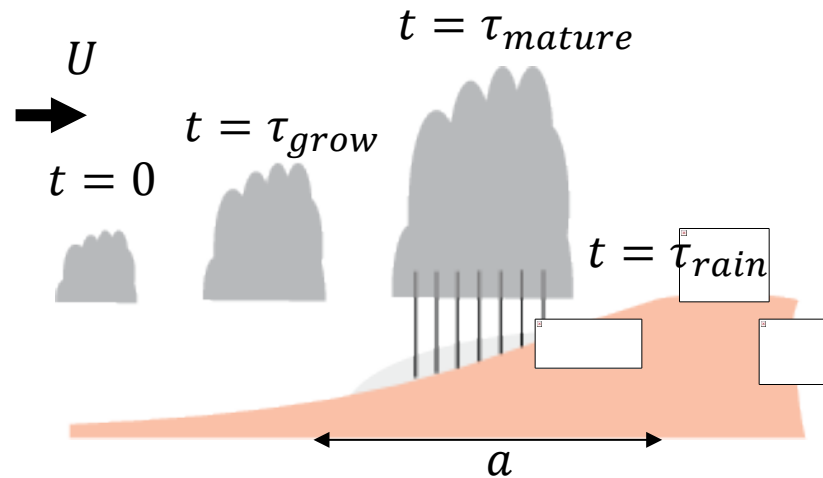
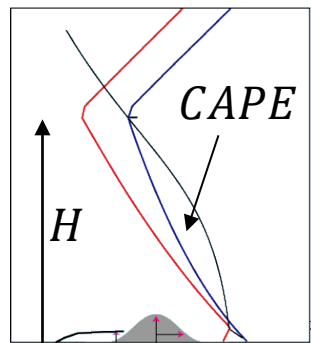
1. Triggering:

$$\frac{h}{LFC} > 1$$

2. Orographic forcing:

$$\frac{h}{a} \left(\frac{w}{U} \propto \frac{h}{a} \right)$$

3. Ratio of advective to convective time scale:



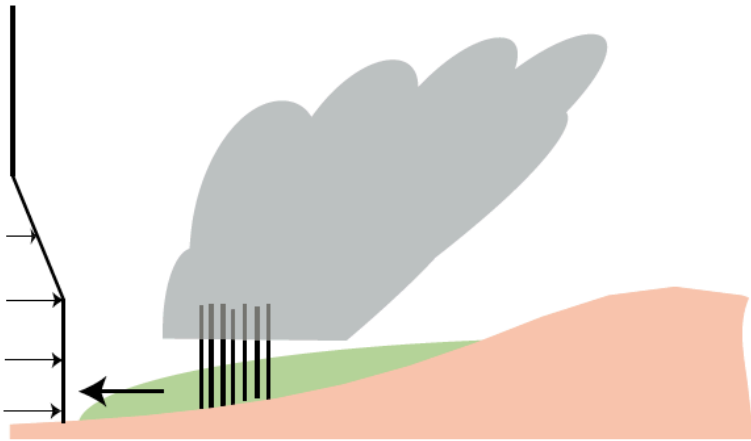
$$\frac{\tau_{adv}}{\tau_{conv}} \propto \frac{a/U}{\tau_{conv}}$$

$$\tau_{conv} = \tau_{grow} + \tau_{mature} + \tau_{rain}$$

Effects of Upstream Wind Variation with Height

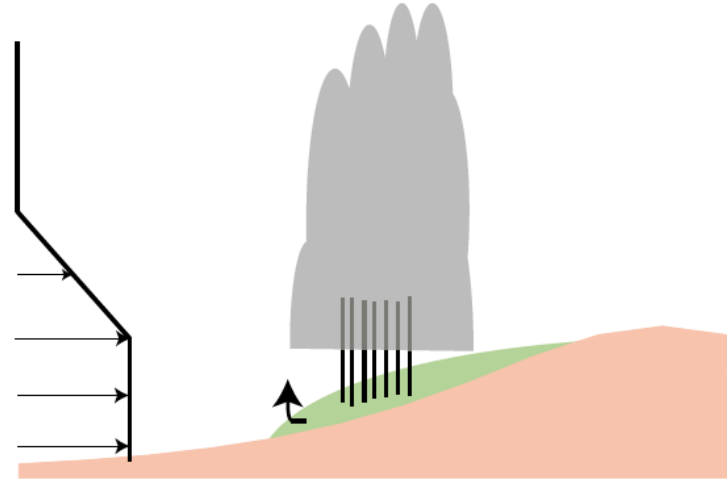
$U(z)$

a) weak low-level wind, zero wind aloft



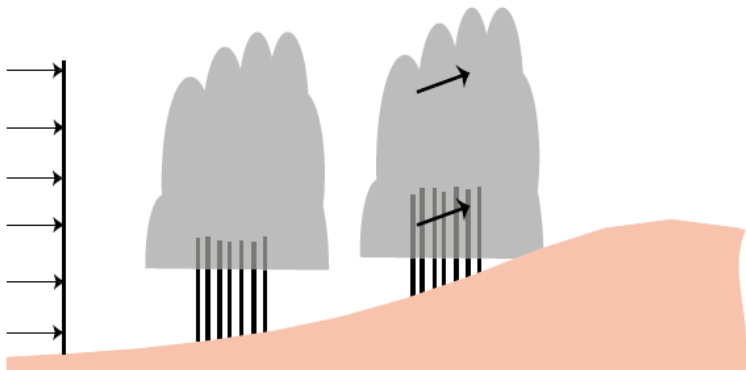
weak advection of clouds aloft --> deep updrafts
--> weak low-level winds allow cold pool to
propagate upstream

b) strong low-level wind, zero wind aloft



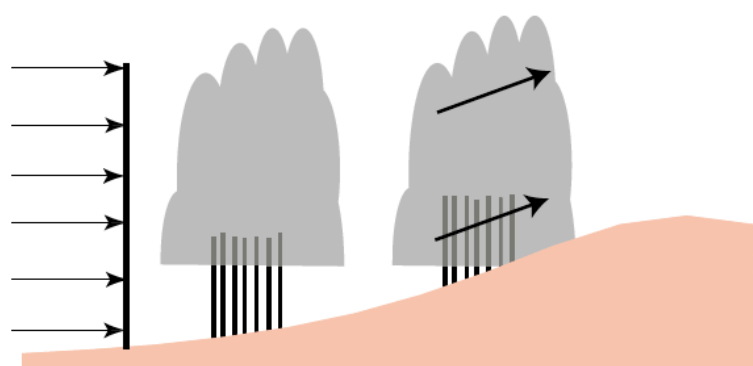
weak advection of clouds aloft --> deep updrafts
--> strong low-level winds counter upstream
cold-pool propagation

c) weak constant wind



advection of clouds --> shallow updrafts --> weak or absent cold pool (MR09 Fig.8)

d) strong constant wind



(Miglietta and Rotunno, 2014 *J Atmos Sci*)

Science Challenges for PMP

- Ensembles of Large Eddy Simulations of Precipitating Clouds for Past Floods with Actual Orography
- Ditto with possible Synoptic/Thermodynamic Environment(s) in Future Climate

Ens. Prob. of Accumul. Rainfall > 5 in. (27 km Neigh.)

Init: 2025-07-04, 0000 UTC
Valid: 2025-07-04, 0600 UTC

